

1 **Synthesis of studies on significant atmospheric electrical effects of major**
2 **nuclear accidents in Chernobyl and Fukushima**

3
4 Accepted version in: [Science of the Total Environment, 733, 139271](#)
5 <https://doi.org/10.1016/j.scitotenv.2020.139271>

6
7 Snežana Dragović^{a*}, Masatoshi Yamauchi^b, Michio Aoyama^c, Mizuo Kajino^d, Jelena
8 Petrović^a, Mirjana Čujić^a, Ranko Dragović^e, Milan Đorđević^e, József Bór^f

9
10 **Abstract**

11 Radioactive materials released during the two most serious nuclear accidents in
12 history, at Chernobyl and Fukushima, caused exceptionally significant contamination and
13 perturbations of the environment. Among them, this paper focuses on the effects related
14 to the atmospheric electricity (AE). Measurements of the most significant disturbances in
15 the values of various AE parameters recorded near ground level are reviewed and the
16 corresponding results are jointly evaluated. The Chernobyl and Fukushima events
17 provided AE characteristics both after long-distance transport (Chernobyl) and short-
18 distance transport including re-suspension (Fukushima). The data indicates that the
19 electrical conductivity of the air is more sensitive to the presence of airborne radioactivity
20 than the atmospheric electric potential gradient (PG). PG, on the other hand, can be
21 monitored more easily and its variation also reflects the vertical redistribution of
22 radionuclides in the air due to their transport, deposition, and re-suspension from the
23 ground. A brief overview of studies on atmospheric transport and deposition of
24 radioactive clouds is given to facilitate the importance of considering the AE
25 measurements in these subjects, and of incorporating those studies in interpreting the
26 results of AE measurements. The AE measurements is particularly important in studying
27 microphysical effects of enhanced radioactivity in the air where no other distance
28 monitoring method exists, both for fair weather conditions wet conditions.

29
30 **Key words:** radionuclides; transport models; atmospheric electric field; potential gradient
31

^aUniversity of Belgrade, Vinča Institute of Nuclear Sciences, POB 522, Belgrade, Serbia
(correspondence: sdragovic@vin.bg.ac.rs)

^bSwedish Institute of Space Physics, Box 812, Kiruna, 98128, Sweden

^cCenter for Research in Isotopes and Environmental Dynamics, Tsukuba University
1-1-1 Tennodai, Tsukuba, Ibaraki 305-8572, Japan

^dMeteorological Research Institute, Japan Meteorological Agency, 1-1
Nagamine, Tsukuba, Ibaraki 305-0052, Japan

^eUniversity of Niš, Faculty of Sciences and Mathematics, Department of Geography, POB 224,
Niš, Serbia

^fResearch Centre for Astronomy and Earth Sciences, GGI, Hungarian Academy of Sciences,
Sopron, Hungary

32 **1. Introduction**

33

34 Nuclear accidents in the Chernobyl Nuclear Power Plant in the Former Soviet
35 Union and in the Fukushima Dai-ichi Nuclear Power Plant in Japan were serious disasters
36 demanding many lives and causing severe damage to the environment. These events were
37 rated at the maximum level (Level 7 - Major Accident) on the International Nuclear and
38 Radiological Event Scale (INES). They caused outstanding disturbances and effects that
39 do not occur normally, so that unambiguously associated response of different natural
40 processes as well as that of the society could be surveyed. The two accidents are similar
41 to each other in a sense that a large amount of radioactive material was injected into the
42 atmosphere. Radionuclides spread and transported mainly by atmospheric circulation also
43 caused severe contamination of the environment at great distances from the places of
44 their origin. Having information on the expectable distribution and various effects of
45 airborne radioactive material is of vital importance from the point of view of making
46 right decisions for minimizing the associated damages. This study summarizes the results
47 obtained from atmospheric electricity (AE) measurements in connection with these two
48 major nuclear accidents.

49 Electrical processes in different regions of the Earth's atmosphere are
50 interconnected and can be considered as elements of one complex system, referred to as
51 the atmospheric global electric circuit (GEC) (Rycroft et al., 2012). The GEC is
52 constantly powered by active thunderstorms and electrified shower clouds around the
53 globe. Charged precipitation and convection currents transport naturally separated
54 electric charges from these clouds to the ground and towards the lower ionosphere,
55 creating a potential difference of 250-300 kV, i.e., an electric field between those two
56 conductive spherical layers. The GEC is closed by low density vertical electric currents
57 flowing over so-called fair weather areas which balance the ongoing charging of this
58 planet-sized capacitor (Haldoupis et al., 2017). The state of GEC is most often
59 characterized by near-surface measurements of the vertical direct current (DC) electric
60 field, the inverse of which is called the atmospheric electric potential gradient (PG)
61 (Nicoll et al., 2019). Direct measurements of electric conductivity, conduction current

62 density, and space charge density of the air contribute to a more complete local
63 characterization of the GEC at different observation sites (Nicoll, 2012).

64 Under steady-state conditions, fair weather conduction current density does not
65 depend on the altitude, therefore it is the vertical profile of the electric conductivity (e.g.,
66 Uman, 1986; Harrison and Bennett, 2007; Rycroft et al., 2008; Williams, 2009), which
67 determines the vertical profile of the electric field (Ogawa and Morita, 1977; Bennett and
68 Harrison, 2008). The electric conductivity is the smallest ($\sim 10^{-14}$ S m⁻¹) at the lowest 10
69 m - 1 km thick layer of the atmosphere, because the ionization rate is the smallest in that
70 layer (Ogawa and Morita, 1977; Tuomi, 1988; Harrison and Bennett, 2007; Rycroft et al.,
71 2008). At the lowest few to few tens of meters from the surface, the major ionization
72 source is ionizing radiation due to the natural radioactivity of the solid Earth, e.g., due to
73 radon gas. At greater altitudes, the ionization is determined mainly by cosmic rays. The
74 flux intensity of cosmic rays, however, decreases exponentially with a rate of one order
75 of magnitude within 15 km altitude difference as they penetrate deep inside the
76 atmosphere from above (Ogawa and Morita, 1977).

77 As a result of low electrical conductivity, the electric potential gradient is the
78 highest in the atmosphere near the ground. This property as well as the ease and relative
79 simplicity of its measurement make PG a sensitive and cost-effective tool for monitoring
80 any change in the lower atmosphere that is manifested in changes of the electric
81 conductivity. For example, the variation of aerosol concentrations in an urban area is
82 unambiguously mirrored by the variation of measured PG values (Harrison and Carslaw,
83 2003; Piper and Bennett, 2012; Jana and Maitra, 2019). Also, the daily variation of the
84 ionized dust through convection (wind) under fair weather conditions is known to cause a
85 corresponding variation in the measured PG values (Kondo, 1959; Harrison, 2003).

86 Radionuclides in the atmosphere increase the local conductivity by contributing to
87 the ionization of the air. Increased local conductivity causes a drop in the PG. This effect
88 was recognized first after nuclear tests in Tucson, Arizona, US (Harris, 1955). The PG
89 measured near the ground dropped suddenly when a rain shower occurred, and this
90 change was also accompanied by an increase of the ion density in the air at the same
91 place. Low PG and the enhancement of the ion density continued for days after the rain
92 was over. That near-surface effect can be attributed rather to the accumulation of the

93 radioactive material on the ground due to washout by precipitation than to radioactive
94 materials suspended in the air. Huzita (1966) discussed that the ionization rate from
95 airborne radioactive dust can be orders of magnitude lower than the ionization rate from
96 natural factors even over 1000 km away from emission source. Therefore, it is rather the
97 radioactive fallout accumulated on the ground which is responsible for the observed
98 change in PG. If the surface is contaminated by artificial radioactive materials, the
99 intensity of the ionizing radiation (beta and gamma rays) exceeds that from natural causes
100 (e.g., radon) by orders of magnitude. This results in a significant increase of electric
101 conductivity and a significant decrease of PG, as illustrated in Figure 1. This is the main
102 reason for radioactive fallout from nuclear tests in the 1950s and 1960s has been
103 diagnosed effectively by atmospheric electricity measurements (Pierce, 1972).

104 Pierce (1959) reported that about 100 km downwind the Windscale nuclear plant,
105 PG dropped statistically by >50% during 1952-1957. This was interpreted as a result of
106 accumulation of radioactive materials, which leaked from the nuclear plant, on the
107 ground. The accumulation may have happened via repeated events of dry deposition and
108 fixation to the soil by independent precipitation. It could also happened due to wet
109 deposition by relatively weak rains. Unfortunately, for Windscale case, no PG-drop
110 events at the arrivals of radioactive dust plumes and rain have been reported, and the
111 increased effect of wet deposition was taken into account only as an average factor.

112 These results also underline the importance of knowing the time evolution of the
113 state and properties of the radioactive contamination injected into the air from the point
114 of view of interpreting AE measurements correctly. The cloud of emitted fission products
115 quickly gets mixed with the air and, depending on the actual meteorological conditions,
116 can be transported far from the source. During atmospheric transport, the ionization
117 potential of the debris cloud is decreasing. This is due to dilution of the radioactive
118 material due to atmospheric mixing, radioactive decay of active isotopes in the plume,
119 and deposition processes (Crandall et al., 1973).

120 Deposition of radionuclides to the ground can be wet or dry (Panitz et al., 1989).
121 Wet deposition is the dominant process during precipitation (rain or snow), while dry
122 deposition is associated with sedimentation of aerosols. Soluble components may
123 undergo a longer atmospheric transport and can act as cloud condensation nuclei

124 (Lelieveld et al., 2012). Considering long distance, radioactive materials are transported
125 by high-altitude (> few km) winds until they meet thick clouds with rain or snow which
126 eventually causes them to fall out to the ground.

127 AE measurements, too, confirmed these long-distance scenarios. PG decreases
128 due to nuclear tests were found to be significant world-wide, affecting even the smallest
129 measured PG values near the ground during 1950s to 1960s (Pierce, 1959, 1972; Kondo,
130 1959; Hamilton and Paren, 1967; Harrison, 2003; Harrison and Ingram, 2005). In these
131 observations, the radioactive fallout was caused by rain (wet deposition) after the world-
132 wide transport in the stratosphere and upper troposphere.

133 In this work, we review the main results of AE-related research triggered by two
134 of the most severe nuclear accidents in the history. One aim of this work is to highlight the
135 specific and unique contribution of AE measurements in assessing the impacts of nuclear
136 accidents on the environment. This contribution contains new knowledge on atmospheric
137 transport processes and effect of radionuclides on AE at the ground level. On the other
138 hand, further investigations are suggested to exploit the results of these measurements
139 mainly by incorporating them into complex studies on the atmospheric transport
140 processes and on the recovering of the contaminated environment. Pullen et al. (2013)
141 pointed out that both accidents revealed the need for improved analytical models for
142 atmospheric but also water dispersion.

143 The timeline of the events and the properties of the corresponding radionuclides
144 emitted into the atmosphere are briefly evoked in section 2. Main methods and the
145 corresponding results of the investigations regarding the transport and the deposition of
146 the radioactive material in the Earth's atmosphere are also summarized in this section.
147 This is included with the intension to aid future research in which AE measurements may
148 be used to support transportation models of airborne material. Additionally,
149 characteristics of radionuclide transport may be considered more in the interpretation of
150 AE measurements. Scientific research on AE-related effects of the two major nuclear
151 accidents is reviewed in section 3. Note that this study focuses intentionally on results
152 which are based on the most significant perturbations caused by the emitted
153 radionuclides. This allows identifying the widest variety of related interactions among
154 environmental parameters. Experiences found at separate measurement sites are jointly

155 evaluated and differences in principal research directions regarding the two cases
156 happened 25 years apart are discussed in section 4. Finally, conclusions are summarized
157 and suggestions regarding further research are given in section 5.

158

159

160 **2. Timeline of the nuclear accidents and studies on the corresponding transport** 161 **and deposition of radionuclides in the atmosphere**

162

163 Radioactive material released to the atmosphere during the accidents can be
164 carried by air masses over long distances before it gets deposited. The distance that
165 radionuclides can travel before they are removed from the air depends on their physical
166 and chemical properties and the weather conditions. In case of nuclear accidents,
167 atmospheric dispersion models play significant role in predicting the movement of
168 released radionuclides in the atmosphere and the deposition patterns of radioactive
169 material all over the world. Comprehensive reviews of the models widely used for
170 characterizing the distribution of radioactive pollutants on local, regional, and global
171 scales are published by Leelóssy et al. (2018) and Bedwell et al. (2018). Uncertainties
172 associated with these models are also addressed in these works. Utilization of
173 atmospheric dispersion models in emergency planning and assessment of the propagation
174 of the hazardous cloud following both Chernobyl and Fukushima accident are discussed
175 by Benamrane et al. (2013).

176

177 **2.1. The accident in Chernobyl**

178

179 The accident at the Chernobyl nuclear power plant located in Ukraine about 20
180 km from its border with Belarus and 130 km north of Kyiv occurred on April 26, 1986
181 during the test of electric control systems at the graphite-moderated light water reactor.
182 Operations conducted during the test led to a rapid increase of reactivity which further
183 induced the vaporization of part of the fuel. The fuel steam expanded rapidly and, finally,
184 caused an explosion which destroyed the largest part of the building. This resulted in the
185 ejection of fuel, core components and some structural parts of the reactor to the

186 surrounding area (UNSCEAR, 2000). The emission of radioactivity from the damaged
187 facility was prolonged by the fire of the graphite moderator which was burning for
188 several days after the explosion. This accidental release caused contamination of both
189 abiotic and biotic components of the environment as it was documented in many
190 investigations following the accident (e.g. Sanada et al., 2002; Korobova et al., 2008;
191 Dragović et al., 2010; Beresford et al., 2016).

192 In the first ten days after the Chernobyl accident, large amounts of radioactive
193 material were released into the atmosphere with the total activities of about 14 EBq
194 which included 1.8 EBq of ^{131}I , 85 PBq of ^{137}Cs and other cesium radioisotopes, 10 PBq
195 of ^{90}Sr and 3 PBq of plutonium radioisotopes (IAEA, 2006). Physical and chemical forms
196 of released material were presented in detail in the report of the International Atomic
197 Energy Agency (IAEA, 2006).

198 The radioactive plume was transported across Europe and two days after the
199 accident measurable radioactive signals were detected from an “unknown” source by
200 automatic monitoring instruments in a Swedish nuclear power plant. Released radioactive
201 materials were pushed into several directions by changing winds, which resulted in an
202 uneven global distribution of radioactive materials and irregular deposited patterns over a
203 wide area. Most affected countries were Belarus, the Russian Federation and Ukraine. A
204 summary of the meteorological conditions (rainfall, changing wind direction) during the
205 Chernobyl accident is given in Atlas map (see De Cort et al., 1998) and IAEA report
206 (IAEA, 2006).

207 Hass et al. (1990) and Pudykiewicz (1988) used EURAD model and simplified
208 version of the Eulerian LRTAP model, respectively, for simulation of the Chernobyl
209 radioactive cloud. Based on chemistry transport model, Hass et al. (1990) show that near-
210 surface part of the radioactive cloud spread out over most of the European countries until
211 May 3, 1986. The model predicted that all other European countries except Ireland,
212 Portugal and major parts of Spain were affected by dry deposition of ^{137}Cs , while
213 Finland, the northern parts of Sweden and Norway, the Alpine region with north Italy and
214 northern Yugoslavia, and the northern parts of Greece and downwind of Chernobyl were
215 affected by wet deposition of ^{137}Cs (Hass et al., 1990). According to findings of
216 Pudykiewicz (1988) the transport of radionuclides occurred in three stages: (1) from 25 to

217 the end of April - radionuclides were transported mostly over Europe, (2) from May 1-7 -
218 radionuclides were transported from the Eastern to Western part of the Northern
219 Hemisphere; and (3) after May 7 - radioactive cloud reached the East and West coasts of
220 North America. Aircraft radioactivity measurements at 5000 m and 10000 m altitude
221 above Europe, the Japan Sea and the West Coast of the United States indicated that
222 accident-derived radioactivity has reached higher levels of the atmosphere than initially
223 estimated.

224 A number of models have been used to assess the atmospheric dispersion of the
225 Chernobyl plume such as MESOS, PATRIC and ADPIC (Ellis, 2003). The Lagrangian
226 puff trajectory model MESOS has been developed with the purpose to simulate the
227 atmospheric transport and dispersion of radionuclides across distances of several hundred
228 kilometers and over (ApSimon et al., 1985). For assessing the period of release, MESOS
229 follows the passage and dilution of a sequence of puffs released at intervals of 3 hours.
230 Meteorological conditions along the paths of the puffs are also taken into account. One of
231 the assumptions of the model is that materials released between these tracked puffs
232 follow intermediate trajectories and dispersion. Daily variations in wind profiles,
233 atmospheric stability, and mixing layer depth were merged over land and, separately,
234 over the sea (ApSimon et al., 1987). Deposition behavior of material on the ground
235 depends on the properties of the radionuclides. The PATRIC model was based on a three-
236 dimensional puff and diffusion model, while the ADPIC model combined mass-
237 consistent wind flow model and a particle-in-cell dispersion model (Ellis, 2003). Due to
238 the rain, hotspots of wet deposition across Europe occurred. Unfortunately, in the mid-
239 eighties of the last century it was impossible to model washout processes, because
240 meteorological precipitation data or fine-scale forecast model precipitation products were
241 not available at that time.

242 Talerko (2005a,b) used the model of atmospheric transport LEDI (Lagrangian-
243 Eulerian Diffusion model) to reconstruct atmospheric transport and ground deposition of
244 ^{137}Cs and ^{131}I , during the first 12 days after the Chernobyl accident, on the Ukrainian
245 territory. A similar pattern was observed for ^{137}Cs and ^{131}I in air and ground
246 contamination fields (Talerko, 2005b). According to Talerko (2005a), the deposition field
247 in central Ukraine was formed predominantly due to dry deposition. Simulations of the

248 transport and deposition of the ^{137}Cs over Europe performed using the coupled model
249 LMDZORINCA showed that ^{137}Cs was spread over very long distances affecting the
250 most of the countries in Europe (Evangelidou et al., 2013b). A comprehensive study of
251 reconstruction of the Chernobyl accident after 30 years was conducted by Evangelidou et
252 al. (2016). Deposition maps of ^{137}Cs for Europe produced based on simulations of the
253 Eulerian transport model LMDZORINCA (INCA) and the Lagrangian particle dispersion
254 model FLEXPART were similar to the Atlas (De Cort et al., 1998). This confirmed the
255 reliability of the models in predicting activity concentrations, cumulative deposition of
256 ^{137}Cs and arrival times of the radioactive fallout (Evangelidou et al., 2016).

257

258 **2.2. The accident in Fukushima**

259

260 The Fukushima Dai-ichi nuclear power plant (FNPP) operated by the Tokyo
261 Electric Power Company (TEPCO) is located in Fukushima Prefecture of the Tohoku
262 region in Japan about 230 km northeast of Tokyo. On March 11, 2011 the earthquake of
263 magnitude 9.0, the largest ever recorded in Japan, occurred as a result of extensive fault
264 ruptures on or near the boundary between the Pacific and North American tectonic plates.
265 The earthquake with a hypocenter depth of 24 km occurred around 130 km offshore the
266 city of Sendai. Immediately after the earthquake, the FNPP reactors were switched off
267 and the fission process was stopped. However, subsequent tsunamis of more than ten
268 meters high inundated the FNPP facilities and disabled the operation of generators for
269 reactor cooling pumps. This led to the melting of nuclear fuel, a series of hydrogen-
270 initiated explosions and severe damage of the cores of three reactors (UNSCEAR, 2013;
271 IAEA, 2015).

272 A large amount of radioactive material was released into the atmosphere from the
273 three damaged cores. Emitted isotopes dispersed across the entire globe in the directions
274 of the prevailing winds (Lin et al., 2015; Thakur et al., 2013) (Figure 2). Meteorological
275 conditions (rainfall, changing wind direction, pressure) during the Fukushima accident
276 are overviewed by the WMO (2013). About two weeks after the accident, radionuclides
277 from Fukushima were detectable across the entire Northern Hemisphere. By April 13,

278 radioactive materials eventually also reached the southern hemisphere in the Asia-Pacific
279 region (CTBTO, 2011).

280 More than 99% of the radioactivity released into the air after Fukushima
281 accident was due to highly volatile radionuclides such as I, Te, Cs, Xe, and Kr (Mathieu
282 et al., 2018). The total released amounts of the top major radionuclides are estimated as
283 ^{133}Xe (6 - 20 EBq), ^{131}I (100 - 400 PBq), ^{132}Te - ^{132}I (87.1 PBq; (Katata et al., 2015)), ^{134}Cs
284 (10 - 35 PBq) and ^{137}Cs (15 - 21 PBq) (Aoyama et al., 2020). The physical half-lives of
285 ^{133}Xe , ^{131}I , ^{132}Te - ^{132}I , ^{134}Cs , and ^{137}Cs are 5.25 d, 8.0 d, 3.2 d-2.3 h, 2.1 y, and 30.0 y,
286 respectively. Chemical and transportation characteristics of fission products from the
287 Fukushima accident were presented in detail in the review paper of Koo et al. (2014).

288 Numerical simulations were commonly used for source term (i.e., quantity of
289 radionuclides released in the atmosphere) estimation following the Fukushima accident.
290 Chino et al. (2011) used an inverse estimation of the source term of iodine and cesium
291 discharged from the FNPP into the atmosphere in the first days after the accident. The
292 method applied was based on the coupling of environmental monitoring data with
293 atmospheric dispersion simulations under the assumption of unit release rate (1 Bq h⁻¹).
294 Several source term assessments have been conducted following the accident using
295 environmental data at both local (Katata et al., 2012; Korsakissok et al., 2013) and large
296 scale (Stohl et al., 2012; Terada et al., 2012; Winiarek et al., 2012).

297 The behavior of the most important radionuclides released after the Fukushima
298 accident has been extensively studied through atmospheric dispersion modelling and
299 model intercomparison studies (e.g., Morino et al., 2011; Stohl et al., 2012; Evangeliou et
300 al., 2013a; Marzo, 2014; SCJ, 2014; Draxler et al., 2015; Katata et al., 2015; Nakajima et
301 al., 2017; Sato et al., 2018). Atmospheric dispersions of radionuclides were simulated by
302 successive uses of the meteorological prediction model PHYSIC and the atmospheric
303 dispersion models PRWDA21 in SPEEDI, and MM5 and GEARN in WSPEEDI-II
304 (Nagai et al., 1999; MEXT, 2007; Terada et al., 2008a,b). Morino et al. (2011) applied a
305 three-dimensional chemical transport model, Models-3 Community Multiscale Air
306 Quality (CMAQ) in order to simulate the distributions of radioactive isotopes ^{131}I and
307 ^{137}Cs at a regional scale during March11-30, 2011. The model roughly reproduced the
308 observed spatiotemporal variations of deposition rates over 15 prefectures in Japan, some

309 discrepancies between the simulated and observed data can be attributed to model
310 uncertainties (Morino et al., 2011).

311 Stohl et al. (2012) estimated that 18% of the total ^{137}Cs deposition until April 20
312 (i.e., 6.4 PBq) occurred over land in Japan. Evangeliou et al. (2013a) used the
313 atmospheric transport model LMDZORINCA to simulate the global transport and
314 deposition of the ^{137}Cs after the Fukushima accident. For the analysis of the atmospheric
315 flows of the ^{137}Cs , Nakajima et al. (2017) used a combination analysis of two aerosol
316 transport models and hourly observed ^{137}Cs concentrations at surface level during March
317 14-23, 2011. The model ensembles revealed the main features of the atmospheric ^{137}Cs
318 distributions and transport routes of eight ^{137}Cs plumes during the studied period.
319 However, some differences between the observed and simulated results were found
320 indicating the need for model improvements (Nakajima et al., 2017).

321 Model inter-comparison (or sensitivity) studies (e.g., Adachi et al., 2013; Morino
322 et al., 2013; Katata et al., 2015; Quérel et al., 2015) can be used to identify the sources of
323 ambiguity in the models. Such studies have been conducted regarding the dispersion and
324 deposition of radionuclides in Japan not only for the sake of simple comparison but also
325 for investigations on transport and deposition mechanisms. The major conclusions
326 obtained from the different inter-comparison studies are coherent. Key points are
327 summarized as follows: (1) there are still significant uncertainties in deposition modeling;
328 (2) it is hard to determine the best model in all aspects, but the ensemble mean can be
329 reasonable in all aspects, and (3) most model inter-comparison studies compare different
330 models with unique or different inputs. The inter-comparison made by Kajino et al.
331 (2019) was different from these in a sense that they compared the output of one model
332 fed with different inputs, i.e., nine different meteorological simulations with a given
333 source term and transport model. This proved to be a useful strategy for evaluating the
334 performance of the applied transport model.

335 Approximately 80% of the radioactive material that was released from FNPP,
336 including what was emitted into the atmosphere, was eventually deposited into the ocean.
337 This enabled a unique approach in this case in which radionuclide measurements in water
338 are used to complete the knowledge on atmospheric circulation and on the matter
339 exchange between the atmosphere and the ocean. East of the coast of the largest Japanese

340 island Honshu is the collision zone of the warm Kuroshio and the cold Oyashio marine
341 currents, which interacted with the air circulation and indirectly influenced the dispersion
342 of radioactive particles and gases into the atmosphere. The behavior of FNPP-released
343 radioactivity in the marine environment are summarized by Aoyama et al. (2016a,b) and
344 Buesseler et al. (2017).

345 North Pacific inventory of ^{137}Cs , calculated from samples from the surface above
346 a standard mixed layer of 50 m over 10 degree x 10 degree areas between 20°N - 50°N,
347 was estimated to be 15-18 PBq (Aoyama et al., 2016a). This result is in accordance with
348 estimates obtained from more comprehensive ocean data sets and by interpolation
349 methods and models - 15 PBq by Inomata et al (2016), and 16 PBq by Tsubono et al.
350 (2016). Close to the FNPP site, direct discharge was the dominant process to form the
351 radioactivity distribution pattern. Inomata et al. (2014) found strong correlations between
352 *in-situ* activities of ^{131}I , ^{134}Cs , and ^{137}Cs measured in surface seawater samples and
353 gamma-ray peak count rates determined by the aerial survey (correlation coefficients
354 were 0.89 for ^{131}I , 0.96 for ^{134}Cs , and 0.92 for ^{137}Cs).

355 The offshore area of high radionuclide activity extended south and southeast from
356 the FNPP. The $^{131}\text{I}/^{137}\text{Cs}$ ratio in surface water of the high-activity area ranged from 0.6
357 to 0.7. Considering the radioactive decay of ^{131}I , they determined that the radionuclides in
358 this area were directly released from FNPP to the ocean. Following the accident, both
359 ^{134}Cs and ^{137}Cs are observed in a wide area in the North Pacific Ocean (Aoyama et al.,
360 2013). The horizontal distribution of ^{134}Cs of FNPP-origin in the western North Pacific
361 Ocean except just in front of the FNPP site showed that the high concentration area was
362 located close to the FNPP accident site. In addition to direct discharge from the FNPP site
363 (Tsunume et al., 2012), deposition from the atmosphere might also have contributed to
364 the increased level of radioactive contamination in this area. This scenario is consistent
365 with the atmospheric transport model study of Honda et al. (2012).

366 Povinec et al. (2013) predicted that radionuclides released in Fukushima accident
367 will reach the west coast of the United States in about five years. The FNPP-derived ^{134}Cs
368 has found to be spread and was first observed at the westernmost station on Line P, an
369 oceanographic sampling line extending 1500 km westward of British Columbia (BC),
370 Canada in June 2012 (Smith et al., 2017). Highly contaminated water reached west coast

371 of North American continent in 2015 (Smith et al., 2017; Kumamoto et al., 2019). Ocean
372 circulation model simulations that are consistent with the time series measurements of
373 Fukushima ^{137}Cs indicate that the 2015–2016 results represent maximum tracer levels
374 (Tsubono et al., 2016). It was also detected in two water samples from the Bering Sea
375 (Kumamoto et al., 2016, 2017). The increase in the Bering Sea was probably derived
376 from the long-range transport of the FNPP1-derived radiocesium into the Bering Sea with
377 the subarctic gyre circulation in the North Pacific (Kumamoto et al., 2019). Huang et al.
378 (2020) reported the presence of FPNN-derived ^{134}Cs and ^{137}Cs in subarctic regions and
379 the Arctic Ocean (Chukchi Sea) in 2017 and attributing it to their transport from the
380 Pacific Ocean into the Bering and Chukchi Seas by ocean currents. They also
381 hypothesized that these radionuclides will be continuously transported into the Canada
382 Basin and Arctic Ocean through the Bering Strait in the next several years.

383

384 **3. Direct and indirect effects of radionuclides from the accidents on near-surface** 385 **AE measurements**

386

387 **3.1. Atmospheric electrical effects of the Chernobyl accident**

388

389 The propagation and extension of the radioactive plume over Europe was
390 confirmed by atmospheric electricity measurements at several measurement sites across
391 the continent.

392

393 **3.1.1. Measurements in Swider, Poland**

394

395 The rich set of measurements in Swider, Poland provided a unique opportunity for
396 the complex analysis of environmental changes related to the event. Recordings of PG
397 were completed by measurements of atmospheric conductivity (for both positive and
398 negative ions), radioactive fallout (Bq m^{-2} , on a weekly cycle), aerosols, i.e., the
399 concentration of condensation nuclei, and meteorological parameters (rain accumulation,
400 wind, and cloudiness). A detailed report on the observations in 1986 was published by
401 Warzecha (1987). Later, the author discussed the extension of the analysis up to 1987

402 (Warzecha, 1991). The arrival of the radioactive contamination at Swider was detected
403 already on the next day of the accident by the simultaneous increase in air conductivity
404 and decrease in PG. The conductivity increased most dramatically in the next couple of
405 days to about ten times the average value before the accident. The lowering of the PG to
406 about 30% of the undisturbed average was less explicit; nevertheless the anti-correlation
407 with the changes in the conductivity was unambiguous. Note that all these changes
408 accompanied an over 100-fold increase in the radioactive fallout density.

409 Several features of the observations were noted. (1) When the ionization intensity
410 was at the peak, the conductivity from negative ions was larger than that from the
411 positive ions. Note that positive conductivity typically exceeds the negative one because
412 negative ions get more easily attached to low mobility large particles (Kawano et al.,
413 1969). (2) High daytime conductivities persisting for several days decreased considerably
414 during nights. It was suggested that this was due to the cutoff of downward transport of
415 radioactive nuclei from higher altitudes because of the stabilization of the boundary layer
416 during nights. (3) Despite the increased level of ionization, the concentration of
417 condensation nuclei near the surface did not grow but it was rather decreased. It was
418 hypothesized that the produced ions were too small to be detected as aerosols. Note that
419 the 1.5 weeks long period following the accident discussed in detail in the report by
420 Warzecha (1987) could be considered as fair weather with practically no precipitation,
421 only few clouds and low wind speeds (Harrison and Nicoll, 2018). In this sense,
422 observation at Swider marks the first-time detection of dry deposition event (Warzecha,
423 1987, 1991).

424 The activity concentration of radioactive fallout peaked in May 1986 and it
425 remained relatively high in the rest of the year. In December 1986, it was still ten times
426 higher than the value before the accident. The anomalously high radioactivity
427 concentration remained for a whole year until only after May, 1987. Measurements of the
428 deposition showed that radioactive substances accumulated in the soil and especially in
429 the grass, where the contamination started to be cleared only in the second half of 1987
430 irregularly. Air conductivities and PG were extremely high mostly in the first three
431 months after the accident until July, after which recovery could be observed so that

432 normal average values returned practically by September-October, 1986 (Warzecha,
433 1987, 1991).

434

435 **3.1.2. Measurements in Helsinki, Finland**

436

437 The atmospheric electricity measurement site at Helsinki Airport is equipped with
438 an electric field probe, a plate for measuring the atmospheric current density, and a pair
439 of aspirators that serve the recording of air conductivity of both positive and negative
440 polarities (Tuomi, 1982, 1988).

441 The arrival of the radioactive contamination at Helsinki Airport caused a gradual
442 increase in conductivities on April 28, i.e., the second day after the accident, due to dry
443 deposition (increase of floating) radioactive substances with the subsiding air currents. At
444 this time the radioactive substances did not stick to the surface. This was indicated by the
445 PG which did not decrease very much. Conductivity values went out of scales of the
446 measurement on April 29 after a precipitation event occurred on-site. The sensitivity of
447 the instrument was adjusted on May 2 when the conductivity was still about ten times
448 higher than its normal level (Figure 3). At the same time, a ten- fold decrease of the PG
449 was registered. Direct measurements of the conduction current were not available during
450 the extreme perturbation, but the current density calculated from the electric field and
451 conductivity, when the latter was available, shows no anomalous deviation in its general
452 level. It was noted, however, that the time series of the current density has become noisier
453 (Tuomi, 1989).

454 The suddenly appearing significant perturbation in the time series of the
455 monitored atmospheric electricity parameters can be attributed to precipitation and the
456 resultant significant wet deposition. The fairly quick decay of the giant peak caused by
457 the perturbation was most probably due to the relatively quick radioactive decay of
458 substances having short half-life (e.g., eight days for ^{131}I). The firm attachment of
459 radionuclides to the ground at Helsinki is consistent with the wet deposition scenario
460 because the majority of the radioactive fallout was cesium (alkali) and iodine (halogen),
461 both of which mingle with soil minerals easily. The attachment to the ground is therefore

462 stronger by wet deposition after high-altitude transport than via dry deposition after low-
463 altitude transport of floating radioactive substances).

464 After the peak, the recovery continued on a slower rate because of radionuclides
465 which have longer half-life (e.g., 31 years for ^{137}Cs vs. eight days for ^{131}I) on one hand,
466 but also due to precipitation which washed the deposited fallout off the surface somewhat
467 deeper into the ground. Conductivity values were recovered to their normal levels
468 practically by August (Tuomi, 1989).

469

470 **3.1.3. Measurements in Uppsala, Sweden**

471

472 Conductivity, PG, and space charge measurements, complemented by recordings
473 of total radioactivity nearby in Uppsala, were running in Marsta Observatory in Sweden
474 in 1986, 1300 km away from the Chernobyl nuclear power plant. In a short publication,
475 Israelsson and Knudsen (1986) reported that the first sign of the arrival of the floating
476 radioactive substance (dry deposition) was a gradual but small increase in air
477 conductivity already on April 27, one day after the accident. The measured parameters
478 changed dramatically when light rain yielding only 0.6 mm precipitation occurred at
479 dawn of April 29. Simultaneously to the ten times increase in background radiation, air
480 conductivity increased to 11 times of its normal levels while PG decreased by a factor
481 seven. In their measurements, the space charge density also decreased by a factor of ten
482 but the authors explained it by limitations of the measurement method due to the
483 disappearing background electric field that is required in the detection method they used.

484 After a period of heavy rainfall during the days May11-13, values of the measured
485 electric parameters returned to about 40% of their normal levels. This was attributed to
486 the rain which could wash much of the depleted radioactive material deeper into the
487 ground (Israelsson and Knudsen, 1986).

488

489 **3.1.4. Measurements in Athens, Greece**

490

491 The set of observations that served the analysis in Athens, inside the city, included
492 PG measurements (at 7 m height above the ground), air conductivity measurements (only

493 positive polarity was available; measured at 4 m height) and small ion concentration
494 (SIC) recordings (both polarities; detected at 4 m height). Simultaneously, activity
495 concentration of several isotopes and the corresponding exposure rates at 1 m height were
496 monitored at 10 km away from the AE measurement (Retalis and Pitta, 1989).

497 The exposure rate at the monitoring site started to increase by the evening of May
498 2, 1986, fully in agreement with predictions from long-range transport models according
499 to which only the substances emitted after April 30 reached the Hellenic region. Parallel
500 to this, air conductivity increased quickly, reaching the maximum of seven times the
501 regular value on May 5. The simultaneous PG drop was only to 53% of its normal value,
502 while the SIC showed four times and five times increase for positive and negative ions,
503 respectively.

504 While activity concentration of the monitored isotopes at 1 m height showed a
505 relatively fast exponential decay rate after May 2, the exposure rate maintained its level
506 after May 5 and started to decrease slowly only after 9-10 days.. This suggests that the
507 altitude distribution of radioactive substances varied in that time period so that a fraction
508 of the radionuclides probably remained floating over the surface for days before they got
509 finally deposited. In fact, the variation of atmospheric electricity parameters, which was
510 somewhat different from the variation of the above-mentioned two parameters of
511 radioactivity, agrees with such scenario. After the effect culminated coherently in the
512 conductivity, PG, and SIC on May 5, a fast recovery was observed until May 8, but these
513 parameters on May 8 still had anomalously different values compared to their regular
514 levels. Their recovery continued, but on a much lower rate which linearly correlated with
515 the logarithm of the activity concentration change of the monitored isotopes (Retalis,
516 1987; Retalis and Pitta, 1989).

517 Note that fair weather conditions were prevalent in Athens during the peak of the
518 effect 3 to 7 May, and the process must be dry deposition. This is also supported by the
519 dynamics of the changes in the AE parameters, which were more gradual and less abrupt
520 compared to variations at the other European observation sites where the wet deposition
521 scenario was unambiguous. The authors called the attention on that the characteristic
522 double peak in the diurnal variation of SIC was completely masked during this period
523 (Retalis and Pitta, 1989).

524

525 **3.2. Atmospheric electrical effects detected after the accident in Fukushima**

526

527 In the Fukushima accident case, long-term, one-minute resolution PG data (raw
528 data sampling rate of 1 Hz at 2.55 m above ground) exists at Kakioka, which is located
529 150 km southwest from the source, and its measurement recovered from the earthquake
530 damage just before the arrival of the radioactive substance (Takeda et al., 2011). The
531 sensor at Kakioka uses Masacart's insulated water-dropper collector that is valid even
532 during rainfall events with response time less than 30 s over 2000 V m^{-1} (Shigeno et al.,
533 2001). As shown in Figure 4 (from Yamauchi et al., 2012), PG became nearly zero twice
534 (green arrows in Figure 4), first after the surface wind from the FNPP brought the
535 radioactive material from the source region without rain, and second after the first rain on
536 March 20-22. The first near-zero period (starting from March 14, 2011) continued more
537 than a day. Without firm fixation by rain during the first PG near-zero period on March
538 14-15, 2011, the contamination forms of the radioactive materials from the FNPP were
539 probably suspended in the near ground air or slightly attached on the soil surface, as
540 illustrated in the middle panel of Figure 5 (from Yamauchi et al., 2012). This means that
541 these radioactive fallouts can easily be re-suspended by the surface wind. If the re-
542 suspension covers all altitudes with the lowest background conductivity (e.g., $< 1 \text{ km}$
543 altitude), the near-surface is no longer the exception for PG altitude distribution (cf.
544 Figure 1) where the finite PG is distributed among these altitudes, leading PG near
545 surface to finite values. In fact the PG increased to nearly the nominal value when strong
546 wind was observed on March 16 (cf. right panel of Figure 5), and then PG fluctuated at
547 values 1/3 to 1/4 of the nominal values after the wind ceased until the rain was recorded
548 (PG dropped to near-zero simultaneously). The observation during this period (from the
549 first near-zero PG period to the first rain) shows that even the dry deposition without rain
550 causes as quick PG drop by one order of magnitude at $> 100 \text{ km}$ distance. The
551 observation also indicates that temporal and partial recovery of the PG must be due to re-
552 suspension of floating radioactive materials from near the surface.

553 The re-suspension of radioactive materials seems to have occurred even after the
554 wet deposition. Such re-suspension is indicated from the diurnal variation of PG (purple

555 arrows in Figure 4), that has different local time (LT) variation and shape from before the
556 accident only until the diurnal variation diminished after April 20. By interpreting the
557 diminished diurnal variation after April 20 as a firm settlement of radioactive fallout
558 (chemical binding with soil minerals), the increase of PG near local noon during this
559 unusual diurnal variation period can be interpreted as wide spread of high-conductance
560 region (re-suspension of radioactive fallout) to substantial altitude, as illustrated in the
561 right panel of Figure 5. The reason that the peak PG values do not reach as large as those
562 of first re-suspension period (March 16-20) is because a substantial amount of surface
563 contamination due to wet deposition (lower panel of Figure 4) makes the near surface still
564 the highest radiation dose rate compared to the higher altitude. Such diurnal re-
565 suspension of the radioactive materials made the surface contamination became leveled,
566 as is seen in the radiation dose rate, which showed faster decay in the heavier
567 contaminated area than in the lighter contaminated area (Yamauchi, 2012; Yamauchi et
568 al., 2012). Such a leveling trend finished late April when the daily variation also
569 diminished. Both indicate the end of substantial re-suspension from these contaminated
570 areas. In other words, the PG data alone indicate that a substantial amount of radioactive
571 particles were floating during the first 50 days after the Fukushima nuclear accident, and
572 that the re-suspension diminished significantly afterward. The resultant time line of
573 floating radionuclide flux is also confirmed by fixed sampling stations at Tsukuba and
574 Takasaki of floating radionuclide (Stoehlker et al., 2011; Kanai, 2012), by sample
575 measurements of floating radioactive particles (Kita et al., 2013), and by in-situ balloon
576 measurements (Fukushima University, 2011).

577 Thus, PG responds differently at the different phases of the transport, fallout, re-
578 suspension, and settlement of the radioactive materials even for the same amount of the
579 radioactive contamination, and the PG data can be used to estimate the motion of floating
580 radionuclides in the air. There was no clear correlation between the time profile of the PG
581 at Kakioka and those of radiation dose rates at the near-by stations (cf. Figure 4). The
582 minor PG changes, that reflect physical processes or dynamics of the local radioactive
583 materials and local ions, could be interpreted by comparing the simultaneous
584 observations of the radiation dose rates and weather records.

585 One last example of such estimate of the change in the radioactive contamination
586 condition is the rain-induced drop of radioactive materials, as indicated by yellow arrows
587 in Figure 4 (April 8 and 18). The nighttime near-zero baseline of the PG at Kakioka
588 slowly recovered during sunny days, but recovery is somewhat interrupted to set back to
589 near zero on these rainy days. Yamauchi et al. (2012) interpreted this setback as fall of
590 radionuclide from tree canopies. The radiation dose rate also shows small spikes on these
591 days, consistent with this scenario. These estimates cannot be made with the radiation
592 dose rate only. Also, high-time resolution monitoring of altitude spread is not easy with
593 the sampling method. In this sense, a combination of PG measurement and radiation dose
594 rate at weather stations is a reliable tool in monitoring the dynamics of the radioactive
595 materials such as the moment of PG drop by dry deposition and re-suspension with high
596 temporal accuracy.

597

598 **3.3. Effects of radionuclides on AE measurements under electrified cloud**

599

600 Using PG measurements alone for diagnosing conductivity enhancements has
601 some limitations: other electric charges, such as cloud electricity, also influence PG in the
602 atmosphere, and mostly the fair weather conditions have been examined until recently.
603 The reason for using PG, despite this problem, to diagnose the electric conductivity is
604 that the electric conductivity is not easy to measure for very low conductivity in the
605 neutral atmosphere. Recently, Yamauchi et al. (2018) found a symptom of PG change
606 even under rain conditions that gives some idea on conductivity and floating
607 radionuclides as described below, although the effect is yet to be quantified.

608 Before the FNPP accident, PG data during the fine weather condition only were
609 analyzed. Data from Kakioka opened new possibility in analyzing the PG data during
610 rain, although the effect of the increased radioactive dust in the atmosphere was identified
611 only in statistics. The difficulty in identifying the effect of increased radiation dose rate in
612 the air under the electrified cloud mainly comes from quick PG changes within a few
613 minutes, as shown in the upper panel of Figure 6 (from Yamauchi et al., 2018), PG
614 decays quickly from its negative peaks corresponding to negative charges that are
615 accumulated at the lower part of the electrified cloud. In the figure, averaged time

616 profiles from the negative PG peak (peak values are between -200 to -400 V m⁻¹) shows
617 shorter time constants for the first 50 days after the Fukushima nuclear accident than the
618 same period of the other years, while no difference was seen before or after this 50-days
619 period. The observed shortening of the time scale is symmetric between rising and decay.

620 The observations can be attributed to the electrostatic shielding effect of the cloud
621 charges by the enhanced ionization as illustrated in the lower panels of Figure 6. For
622 positive peaks, no notable difference was detected on the time scale to/from the peak
623 between 2011 and the other years. This is consistent with the concentration of the
624 radioactive particle at low altitude (< 500 m according to in-situ gamma ray observations
625 by radiosonde), where negative cloud charges exceed the positive charges. The PG data
626 under electrified cloud gives additional information on the altitude range of the re-
627 suspended radioactive fallouts. The results open up a new possibility to use PG as an
628 independent monitor of radioactivity at some altitudes even during long period of
629 cloudy/rainy conditions.

630

631 **3.4. Indirect effects of radionuclides on AE measurements**

632

633 Radionuclides released to the atmosphere get attached to aerosols which can
634 further influence different atmospheric processes, and transport dynamics of radioactive
635 aerosols can be different from the normal case. Aerodynamic size and hygroscopicity are
636 the two most important microphysical properties of aerosols, which determine
637 atmospheric lifetime and deposition amounts. The modeling studies following Fukushima
638 accident pointed out that knowledge of these aerosol properties is just as important as the
639 accuracy of the meteorological simulations and emission scenarios (Saito et al., 2015).
640 Radioactive particles influence charge transfer processes within the clouds through ion-
641 induced aerosol particle formation (Williams and Mareev, 2014). In fact, enhanced
642 thunderstorm activity and higher occurrence of lightning in Sweden after the Chernobyl
643 accident can be explained by the increased radioactivity (Israelsson et al., 1987), although
644 the observed increase of 30% may simply be due to annual anomaly. The influences may
645 thus be diverse and are supposed to depend on a large number of factors, e.g., physical
646 and chemical characteristics of radionuclides, size and density of the radioactive plume,

647 distribution of atmospheric pressure, general atmospheric circulation, local winds,
648 weather conditions at the time of the accident, altitude of the emitting source, regional
649 ocean currents, thermohaline circulation, and regional as well as global climate
650 anomalies.

651

652 **4. Discussion**

653

654 Recognizing the need to improve the understanding of processes governing
655 radionuclide behavior in the environment, both accidents led to the revival of research on
656 preparing for similar scenarios in the future. Following the Chernobyl accident, a number
657 of decision-supporting models have been developed. These models incorporated
658 information on factors which influence the concentration of airborne radionuclides.
659 Meteorological data, dry and wet deposition processes and orographic data were
660 considered. In the decades that followed, the focus was more on dispersion models which
661 vary in terms of their structure, specific output, spatial scale and complexity. They aimed
662 at providing predictions for atmospheric dispersion to support emergency response in
663 accidental releases. Such predictions of atmospheric dispersion assuming point unit
664 source were actually issued on the web by different teams after the accident in
665 Fukushima, e.g., by Norway, Austria, France, and Germany, but measurement-based
666 open warning was not provided due to the lack of sufficient active monitoring stations,
667 particularly for the floating radioactive materials in the air.

668 Airborne radionuclides from the considered nuclear accidents significantly
669 perturbed the electrical characteristics of air and thus contributed significantly to research
670 on the atmospheric electricity. AE-related effects of the Chernobyl accident were
671 reported from a very wide region even from over 1000 km distance in countries in North-
672 East, East, and South-East Europe. A variety of atmospheric electricity measurements
673 were running at the different monitoring stations. These monitoring capabilities include
674 PG and air conductivity measurements at all stations, with additional complete
675 observations of space charge and atmospheric current density, aerosols, condensation
676 nuclei and small ion concentration at some stations. Being aware of the importance of wet
677 deposition of airborne radioactive material, various sets of meteorological data were

678 incorporated in studies. Nuclear-accident-related studies also considered various
679 reference measurements of exposure rates or environmental radioactivity (usually from
680 the nearest monitoring site) in the form of fallout activity concentration at ground or near
681 the surface in the air.

682 Such rich set of measurements allowed the examination of the response of
683 various parameters of atmospheric electricity to the Chernobyl-origin nuclear substances
684 that arrived at the region in the air. Air conductivity was found to be more sensitive to
685 radioactive contamination than the PG at most sites. The ratio of the change rate in air
686 conductivity to that in PG were 3.0 (10/3.33) in Poland, 1.57 (11/7) in Sweden, 3.7
687 (7/1.89) in Greece, and 1.0 (10/10) in Finland. On one hand, variation of this ratio can be
688 attributed to local factors affecting the sensitivity of the measurements. For example,
689 conductivity and PG measurements in Athens were made at different heights. On the
690 other hand, the vertical distribution of the radionuclides in the air and the type of the
691 deposition (wet or dry) also has an influence on this ratio. While all measured
692 atmospheric electricity parameters were affected by the perturbation caused by the
693 elevated level of radioactivity in the air, the effect on the electric current density in the air
694 was found to be very small (only some increased noise was measured in Finland). This
695 suggests that the large-scale GEC was not affected despite the dramatic changes in the
696 local electric properties of the environment.

697 In addition to the observation of fundamental effects, i.e., increasing conductivity
698 and decreasing PG, several observations were made after the Chernobyl accident on the
699 interconnections within a larger set of atmospheric electricity parameters, aerosol as well
700 as ion concentration variations. However, perturbations manifested in specific variations
701 in the composition and size of ions and aerosols could be observed at limited sites where
702 the corresponding measurements were available. This resulted in that some observations
703 could not be confirmed by independent measurements. Such unconfirmed observations
704 include the anomalously higher conductivity of negative ions compared to that of positive
705 ones in Poland, the unexpected decrease in space charge density recorded in Sweden, or
706 the characteristic change in the diurnal variation of small ion concentration detected in
707 Greece. On the other hand, the hypothesis on the observed lowered concentration of
708 condensation nuclei in Poland (i.e., that mostly small ions were produced) was indirectly

709 supported by the measured (and slightly asymmetric) growth of positive and negative
710 small ion concentrations noted in Greece. These experiences underline the importance of
711 running as many different measurements at an observation site as possible.

712 For Fukushima accident, such multi-point comparison was not possible partly
713 because the radionuclides spread mainly over ocean where no observatory exists, and
714 Kakioka without conductivity measurements was the only PG station that was active
715 within the affected area. On the other hand, there was a dense network of the
716 meteorological monitoring sites. Therefore, atmospheric electricity-related studies on the
717 Fukushima accident concentrated only on the PG and discuss its local variation in detail
718 together with the variation of meteorological parameters (i.e., rain and wind) as well as
719 radiation dose rate from a nearby monitoring site. The high time resolution (10 min)
720 meteorological records and radiation dose rate allowed making detailed studies of the
721 changing effect of dry deposition for several days. In these studies, the emphasis was
722 placed on identifying as many factors as possible from those which drive the variations of
723 PG in association with the environmental changes caused by the arriving nuclear
724 contamination in the air. Dry deposition scenarios were present in Poland and in Greece,
725 too, after the Chernobyl accident, but the discussion in the corresponding papers focused
726 more on the responses of the various measured atmospheric electricity parameters to the
727 radioactive material and not very specifically behaviors of arrived radioactive material
728 such as re-suspension and its relation to meteorological parameters. In studies associated
729 with the accident in Chernobyl, the primary role of meteorological data was to confirm
730 fair weather conditions and precipitation events.

731 The analyses based on PG measurements in Kakioka could benefit from the
732 sequence of rain events causing wet deposition after the days of dry deposition/re-
733 suspension period. In this way, both dry and wet deposition scenarios were studied.
734 Furthermore, partial re-suspension of the wet-deposited radioactive material could be
735 recognized. Note that a similar sequence of dry and wet deposition scenarios also
736 happened in Sweden and Finland after the Chernobyl accident but the length of the dry
737 deposition periods was only one-two days in both cases and detailed analysis on the PG
738 variations in connection with the changing weather conditions was not published.
739 Correspondences between PG variations and meteorological conditions, transport,

740 settlement, and fallout processes are known much better nowadays as a result of
741 extensive studies in connection with the Fukushima event.

742

743 **5. Conclusions**

744

745 Contribution of studies on the electrical environment to know the time evolution
746 of the state and properties of radioactive contamination, which has primary importance
747 for predicting, estimating, and surveying the consequences of large-scale emission events
748 of radioactive substance, were reviewed in this study. Main findings are found after
749 Chernobyl and Fukushima nuclear accidents. AE-related studies published in connection
750 with the Chernobyl accident compared the temporal variations of various, locally
751 measured AE parameters either to one-another and to local measurements of radioactivity
752 when those were available. While effects of radionuclides from Fukushima on the
753 electrical properties of the atmosphere have been investigated by focusing on PG
754 measurements, meteorological conditions were more carefully considered than in earlier
755 studies and the potential of PG measurements was demonstrated in differentiating
756 between events of dry and wet deposition as well as re-suspension of the airborne
757 radioactive material.

758 According to measurements after these accidents, air conductivity seems to be the
759 most sensitive AE parameter from the point of view of detecting the local effects of
760 airborne radionuclides, whereas the atmospheric electric current density was fairly stable
761 during the episodes of radioactive contamination. This suggests that, PG can be used to
762 indicate the changes in air conductivity through Ohm's law, and that the overall state of
763 the GEC is not changed significantly (at least on short terms) because of the intake of
764 radionuclides in the atmosphere. Despite this finding, the ratio between the change rates
765 of air conductivity and PG varied significantly (1.0-3.7) at different observation sites.
766 Characteristics of the local measuring environment may contribute to this variation, the
767 origin of which remains to be identified yet.

768 According to the experience from Athens after the Chernobyl accident, also the
769 concentration of small ions is sensitive to the appearance of airborne radionuclides, but
770 measuring it is not easy compared to the monitoring of PG. Space charge measurements,

771 as they were implemented in Sweden, could not indicate radioactivity-related effects
772 unambiguously. Similarly, variations of the aerosol concentration alone are not suitable
773 for detecting the arrival of radioactive contamination at the monitoring site. These
774 measurements must be evaluated together with records of proven indicator parameters.

775 Studies in connection with the Fukushima event clearly demonstrated that PG is a
776 very useful and convenient parameter for studying the effects of nuclear contamination in
777 the atmosphere. Not only it is apparently easy to measure, but its variation was found to
778 mirror the characteristics of vertical transport processes and the changes in the local
779 distribution of the radioactive material, for which no other remote sensing method is
780 found. PG can be used to distinguish between wet and dry deposition scenarios and to
781 detect re-suspension of radionuclides from the ground (even after wet deposition).
782 Perturbations and smoothing of the diurnal PG variation refer to redistributions of the
783 radioactive material in the environment corresponding to its falling from tree canopies or
784 altitude spreading of the contaminated air layer, respectively. Thus, PG measurement,
785 together with radiation dose rate measurement, constructs a unique method to monitor
786 floating radionuclides at more than 10 m altitude, which is not possible to measure by
787 other ground-based methods. Therefore, PG stations are a good infrastructure to deploy
788 around nuclear facilities.

789 The dynamics of the changes registered in PG or in air conductivity can be used to
790 infer variations in the composition and amount of radioactive material deposited on the
791 surface of the ground. The ratio of radioactive components with different half-time can be
792 inferred from the recovery rate of those AE parameters in the local environment. On
793 longer terms, recovery and clearing of the ground from radionuclides can be monitored
794 the same way. Note that the change reflected by the variation of AE parameters was
795 found to be different from what had been obtained by direct measurements of near-
796 surface exposure and activity concentration.

797 These experiences indicate well that AE measurements could be used to verify and
798 support models on atmospheric transport, distribution, and the composition of clouds of
799 radionuclides. On the other hand, the corresponding direct isotope measurements and
800 model outputs can be used to interpret the details of the observed variations of AE
801 parameters. Some of the most widely used transport and dispersion models are briefly

802 reviewed in section 2 of this study for consideration in planning future research on the
803 subject.

804 A unique benefit of AE measurement is that they can serve studies on microphysical
805 effects of airborne radioactivity. Aerosol and small ion measurements and detection of
806 space charge density at some of the observation sites revealed asymmetry in ionization
807 and attachment processes in terms of the polarity of the involved particles in response to
808 increased radioactivity in air. Several corresponding results need to be confirmed yet
809 either by modeling or by other measurements. In any case, these results emphasize that it
810 is worth equipping atmospheric electricity monitoring sites with a more complete set of
811 instrumentation. Measuring atmospheric conductivity in addition to PG and a well-
812 equipped meteorological station is particularly suggested.

813 Results on the responses of atmospheric electricity parameters to nuclear
814 substances carried with the air in case of both Chernobyl and Fukushima demonstrate
815 well the high potential of atmospheric electricity measurements in detecting and
816 monitoring the presence of radioactive material in the atmosphere. A novel approach
817 indicates that PG measurements are capable of fulfilling this task not only in fair weather
818 but also in the presence of clouds.

819

820 **Acknowledgements**

821 The authors SD, JP, MĆ, RD, MĐ and JB would like to acknowledge networking
822 support by the COST Action CA15211. The authors SD, JP, MĆ, RD and MĐ
823 acknowledge the support of the Ministry of Education, Science and Technological
824 Development of the Republic of Serbia (project III43009). MK was supported by the
825 Japanese Radioactivity Survey from the NRA, Japan, the Japan Society of the Promotion
826 of Science (JSPS) and Ministère des Affaires Étrangères et du Développement
827 International (MAEDI) under the Japan - France Integrated Action Program (SAKURA),
828 and the Environmental Research and Technology Development Fund of the
829 Environmental Restoration and Conservation Agency (ERCA) (1-1802). MA was
830 supported in part by the radioactive survey and the research fund of the Ministry of
831 Education, Culture, Sports, Science and Technology (MEXT), Japan (Houshanou-
832 chousa-kenkyuhi, FY2011–2014) and in part by the J-RAPID fund of the Japan Science
833 and Technology Agency for the project entitled “Investigation and Prediction of Impacts
834 of the 2011 off the Pacific coast of Tohoku Earthquake on Marine Environment,
835 FY2011–2012”. MA was also supported by research fund of Institute of Environmental
836 Radioactivity, Fukushima University, during the period from FY2014 to FY2018, too.
837 Contribution of JB was supported by the National Research, Development and Innovation
838 Office, Hungary-NKFIH, K115836. The atmospheric electricity (PG) data at Kakioka and

839 relevant weather data are provided by the Japan Meteorological Agency and publicly
840 available at <http://www.kakioka-jma.go.jp/obsdata/metadata/en/> and
841 <http://www.data.jma.go.jp/gmd/risk/obsdl/index.php>. The radiation dose rate data is
842 provided by Ibaraki-ken prefecture.
843

844 References

- 845 Adachi, K., Kajino, M., Zaizen, Y., Igarashi, Y., 2013. Emission of spherical cesium-
846 bearing particles from an early stage of the Fukushima nuclear accident. *Sci. Rep.*
847 3: 2554. <https://doi.org/10.1038/srep02554>.
- 848 Aoyama, M., Tsumune, D., Inomata, Y., Tateda, Y., 2020. Mass balance and latest fluxes
849 of radiocesium derived from the Fukushima accident in the western North Pacific
850 Ocean and coastal regions of Japan, *J. Environ. Radioact.* 217, 106206.
851 <https://doi.org/10.1016/j.jenvrad.2020.106206>.
- 852 Aoyama, M., Hamajima, Y., Hult, M., Oka, E., Tsumune, D., Kumamoto, Y., 2016a. ^{134}Cs
853 and ^{137}Cs in the North Pacific Ocean derived from the March 2011 TEPCO
854 Fukushima Dai-ichi Nuclear Power Plant accident, Japan. Part one: surface
855 pathway and vertical distributions, *J. Oceanogr.* 72, 53-65. [doi: 10.1007/s10872-](https://doi.org/10.1007/s10872-015-0335-z)
856 [015-0335-z](https://doi.org/10.1007/s10872-015-0335-z).
- 857 Aoyama, M., Kajino, M., Tanaka, T.Y., Sekiyama, T.T., Tsumune, D., Tsubono, T.,
858 Hamajima, Y., Inomata, Y., Gamo, T., 2016b. ^{134}Cs and ^{137}Cs in the North Pacific
859 Ocean derived from the March 2011 TEPCO Fukushima Dai-ichi Nuclear Power
860 Plant accident, Japan. Part two: estimation of ^{134}Cs and ^{137}Cs inventories in the
861 North Pacific Ocean, *J. Oceanogr.* 72, 67-76. [https://doi.org/10.1007/s10872-015-](https://doi.org/10.1007/s10872-015-0332-2)
862 [0332-2](https://doi.org/10.1007/s10872-015-0332-2).
- 863 Aoyama, M., Uematsu, M., Tsumune, D., Hamayima, Y., 2013. Surface pathway of
864 radioactive plume of TEPCO Fukushima NPP1 released ^{134}Cs and ^{137}Cs ,
865 *Biogeosciences*, 10, 3067-3078. <https://doi.org/10.5194/bg-10-3067-2013>.
- 866 ApSimon H.M., Wilson, J.J.N., 1987. Modelling atmospheric dispersal of the Chernobyl
867 release across Europe. *Boundary Layer Meteor.* 41, 123-133.
868 <https://doi.org/10.1007/BF00120435>.
- 869 ApSimon, H.M., Goddard A.J.H., Wrigley J., 1985. Long-range atmospheric dispersion
870 of radioisotopes – i. the MESOS model. *Atmos. Environ.* 19, 99-111.
871 [https://doi.org/10.1016/0004-6981\(85\)90141-6](https://doi.org/10.1016/0004-6981(85)90141-6).
- 872 Bedwell, P., Wellings, J., Leadbetter, S., Tomas, J., Andronopoulos, S., Korsakissok, I.,
873 Périllat, R., Mathieu, A., Geertsema, G., Klein, H., de Vries, H., Hamburger, T.,
874 Pázmándi, T., Rudas, C., Sogachev, A., Szántó, P., 2018. Guidelines detailing the
875 range and distribution of atmospheric dispersion model input parameter
876 uncertainties. In: *Guidelines ranking uncertainties for atmospheric dispersion*
877 *European joint programme for the integration of radiation protection research*.
- 878 Benamrane, Y., Wybo, J.-L., Armand, P., 2013. Chernobyl and Fukushima nuclear
879 accidents: what has changed in the use of atmospheric dispersion modeling? *J.*
880 *Environ. Radioact.* 126, 239-252. <https://doi.org/10.1016/j.jenvrad.2013.07.009>.
- 881 Bennett, A.J., Harrison, R.G., 2008, Variability in surface atmospheric electric field
882 measurements. *J. Phys. Conf. Ser.* 142, 012046IOP. [https://doi.org/10.1088/1742-](https://doi.org/10.1088/1742-6596/142/1/012046)
883 [6596/142/1/012046](https://doi.org/10.1088/1742-6596/142/1/012046).

884 Beresford N., Fesenko S, Konoplev A, Skuterud L, Smith J.T., Voigt G, 2016. Thirty
885 years after the Chernobyl accident: What lessons have we learnt? *J. Environ.*
886 *Radioact.* 157, 77-89. <https://doi.org/10.1016/j.jenvrad.2016.02.003>.

887 Bonjean, F., Lagerloef, G.S.E., 2002. Diagnostic model and analysis of the surface
888 currents in the tropical Pacific Ocean. *J. Phys. Oceanogr.* 32, 2938–2954.
889 [https://doi.org/10.1175/1520-0485\(2002\)032<2938:DMAAOT>2.0.CO;2](https://doi.org/10.1175/1520-0485(2002)032<2938:DMAAOT>2.0.CO;2).

890 Buesseler, K., Aoyama, M., Benitez-Nelson, K., Charmasson, S., Higley, K., Maderich,
891 V., Masqué, P., Morris, P.J., Oughton, D., Smith, J.N., 2017. Fukushima Daiichi–
892 Derived Radionuclides in the Ocean: Transport, Fate, and Impacts. *Ann. Rev.*
893 *Mar. Sci.* 9, 173-203. <https://doi.org/10.1146/annurev-marine-010816-060733>.

894 Chang, C.-C., Burr, G.S., Jull, A.J.T., Russell, J., Priyadarshi, A., Lin, M., Thiemens, M.,
895 Biddulph, D., 2019. Measurements of ¹²⁹I in the Pacific Ocean at Scripps Pier and
896 Pacific Northwest sites: A search for effects from the 2011 Fukushima Daiichi
897 Nuclear Power Plant accident and Hanford. *Sci. Total Environ.* 689, 1023-1029.
898 <https://doi.org/10.1016/j.scitotenv.2019.06.372>.

899 Chino, M., Nakayama, H., Nagai, H., Terada, H., Katata, G., Yamazawa, H., 2011.
900 Preliminary estimation of release amounts of ¹³¹I and ¹³⁷Cs accidentally
901 discharged from the Fukushima Daiichi nuclear power plant into the atmosphere.
902 *J. Nucl. Sci. Technol.* 48, 1129-1134.
903 <https://doi.org/10.1080/18811248.2011.9711799>.

904 Crandall, W.K., Molenkamp, C.R., Williams, A.L., Fulk, M.M., Lange, R., Knox, J.B..
905 1973. An Investigation of scavenging of radioactivity from nuclear debris clouds:
906 research in progress, Lawrence Livermore Laboratory Report, UCRL-51328,
907 <https://inis.iaea.org/collection/NCLCollectionStore/Public/04/068/4068239.pdf>.

908 CTBTO (Comprehensive Nuclear-Test-Ban Treaty Organization), 2011. Fukushima
909 Related Measurements by the CTBTO. [Online] Available at:
910 [http://www.ctbto.org/press-centre/highlights/2011/fukushima-related-](http://www.ctbto.org/press-centre/highlights/2011/fukushima-related-measurements-bythe-ctbto/fukushima-related-measurements-by-the-ctbto-page-1/)
911 [measurements-bythe-ctbto/fukushima-related-measurements-by-the-ctbto-page-1/](http://www.ctbto.org/press-centre/highlights/2011/fukushima-related-measurements-bythe-ctbto/fukushima-related-measurements-by-the-ctbto-page-1/).

912 De Cort, M., Dubois, G., Fridman, S.D., Germenchuk, M.G., Izrael, Y.A., et al., 1998.
913 Atlas of Caesium Deposition on Europe after the Chernobyl Accident. Office for
914 Official Publications of the European Communities, Luxembourg. ISBN 92-828-
915 3140-X.

916 Dragović, S., Howard, B.J., Caborn, J.A., Barnett, C.L., Mihailović, N., 2010. Transfer of
917 natural and anthropogenic radionuclides to ants, bryophytes and lichen in a semi-
918 natural ecosystem. *Environ. Monit. Assess.* 166, 677-686.
919 <https://doi.org/10.1007/s10661-009-1032-4>.

920 Draxler, R., Arnold, D., Chino, M., Galmarini, S., Hort, M., Jones, A., Leadbetter, S.,
921 Malo, A., Maurer, C., Rolph, G., Saito, K., Servranckx, R., Shimbori, T., Solazzo,
922 E., Wotawa, G., 2015. World Meteorological Organization’s model simulations of
923 the radionuclide dispersion and deposition from the Fukushima Daiichi nuclear
924 power plant accident. *J. Environ. Radioact.* 139, 172-184.
925 <https://doi.org/10.1016/j.jenvrad.2013.09.014>.

926 Ellis J., 2003. Use of atmospheric models in response to the Chernobyl disaster, in:
927 Tracking and Predicting the Atmospheric Dispersion of Hazardous Material
928 Releases: Implications for Homeland Security. Committee on the Atmospheric

929 Dispersion of Hazardous Material Releases, National Research Council, pp. 87-
930 88.

931 Evangeliou, N., Balkanski, Y., Cozic, A., Møller A.P., 2013a. Global transport and
932 deposition of ¹³⁷Cs following the Fukushima Nuclear Power Plant Accident in
933 Japan: Emphasis on Europe and Asia using high-resolution model versions and
934 radiological impact assessment of the human population and the environment
935 using interactive tools. *Environ. Sci. Technol.* 47, 5803-5812.
936 <https://doi.org/10.1021/es400372u>.

937 Evangeliou, N., Balkanski, Y., Cozic, A., Møller, A.P., 2013b. Simulations of the
938 transport and deposition of ¹³⁷Cs over Europe after the Chernobyl Nuclear Power
939 Plant accident: influence of varying emission-altitude and model horizontal and
940 vertical resolution. *Atmos. Chem. Phys.* 13, 7183-7198.
941 <https://doi.org/10.5194/acp-13-7183-2013>.

942 Evangeliou, N., Hamburger, T., Talerko, N., Zibtsev, S., Bondar, Y., Stohl, A.,
943 Balkanski, Y., Mousseau, T.A., Møller, A.P., 2016. Reconstructing the Chernobyl
944 Nuclear Power Plant (CNPP) accident 30 years after. A unique database of air
945 concentration and deposition measurements over Europe. *Environ. Pollut.* 216,
946 408-418. <https://doi.org/10.1016/j.envpol.2016.05.030>.

947 Fukushima University (2011) press release, 11 May, 2011. [http://www.sss.fukushima-](http://www.sss.fukushima-u.ac.jp/sonde\data)
948 [u.ac.jp/sonde\data](http://www.sss.fukushima-u.ac.jp/sonde\data).

949 Haldoupis, C., Rycroft, M. Williams, E., Price, C., 2017. Is the “Earth-ionosphere
950 capacitor” a valid component in the atmospheric global electric circuit? *J. Atmos.*
951 *Sol-Terr. Phys.*, 164, 127-131. <https://doi.org/10.1016/j.jastp.2017.08.012>.

952 Hamilton, R.A., Paren, J.G., 1967. The influence of radioactive fallout on the
953 atmospheric potential gradient. *Meteorol. Mag.* 96, 81-85.

954 Harris, D.L., 1955. Effects of radioactive debris from nuclear explosions on the electrical
955 conductivity of the lower atmosphere. *J. Geophys. Res.* 60, 45-52.
956 <https://doi.org/10.1029/JZ060i001p00045>.

957 Harrison R.G., Nicoll K.A., 2018. Fair weather criteria for atmospheric electricity
958 measurements. *J. Atmos. Sol. Terr. Phys.* 179, 239-250.
959 <https://doi.org/10.1016/j.jastp.2019.01.003>.

960 Harrison, R. G., Carslaw, K. S., 2003. Ion-aerosol-cloud processes in the lower
961 atmosphere. *Rev. Geophys.* 41. <https://doi.org/10.1029/2002RG000114>.

962 Harrison, R.G., 2003. Twentieth century atmospheric electrical measurements at the
963 observatories of Kew, Eskdalemuir and Lerwick. *Weather* 58, 11-19.
964 <https://doi.org/10.1256/wea.239.01>.

965 Harrison, R.G., Bennett, A.J., 2007. Cosmic ray and air conductivity profiles retrieved
966 from early twentieth century balloon soundings of the lower troposphere. *J.*
967 *Atmos. Sol. Terr. Phys.* 69, 515–527. [10.1016/j.jastp.2006.09.008](https://doi.org/10.1016/j.jastp.2006.09.008).

968 Harrison, R.G., Ingram, W.J., 2005. Air-earth current measurements at Kew, London,
969 1909-1979. *Atmos. Res.* 76, 49-64.
970 <https://doi.org/10.1016/j.atmosres.2004.11.022>.

971 Hass, H., Memmesheimer, M., Geiß, H., Jakobs, H.J., Laube, M., Ebel, A., 1990.
972 Simulation of the Chernobyl radioactive cloud over Europe using the Eurad
973 model. *Atmos. Environ.* 24, 673-692. [https://doi.org/10.1016/0960-](https://doi.org/10.1016/0960-1686(90)90022-F)
974 [1686\(90\)90022-F](https://doi.org/10.1016/0960-1686(90)90022-F).

- 975 Honda, M.C, Aono, T., Aoyama, M., Hamajima, Y, Kawamaki, H., Kitamura, M.,
 976 Masumoto, Y., Miyazawa, Y., Takigawa, M., Saino, T., 2012. Dispersion of
 977 artificial caesium-134 and-137 in the western North Pacific one month after the
 978 Fukushima accident. *Geochem. J.* 46, e1-e9.
 979 <https://doi.org/10.2343/geochemj.1.0152>.
- 980 Huang, D., Lin, J., Du, J., Yu, T., 2020. The detection of Fukushima-derived radiocesium
 981 in the Bering Sea and Arctic Ocean six years after the nuclear accident. *Environ.*
 982 *Pollut.* 256, 113386, <https://doi.org/10.1016/j.envpol.2019.113386>. Huzita A.,
 983 1966. Effect of Radioactive Fallout Upon the Electrical Conductivity of the Lower
 984 Atmosphere. *J. Geomagn. Geoelectr.* 18, 363-372.
 985 <https://doi.org/10.5636/jgg.18.363>.
- 986 IAEA (International Atomic Energy Agency), 2006. Environmental consequences of the
 987 Chernobyl accident and their remediation: twenty years of experience / report of
 988 the Chernobyl, Forum Expert Group “Environment”, IAEA, Vienna, Austria,
 989 ISBN 92-0-114705-8.
- 990 IAEA (International Atomic Energy Agency), 2015. The Fukushima Daiichi Accident,
 991 IAEA, Vienna, Austria, ISBN 978-92-0-107015-9.
- 992 IEM (Iowa Environmental Mesonet),
 993 http://mesonet.agron.iastate.edu/sites/windrose.phtml?network=JP_ASOS&station=RJSS.
 994
- 995 Inomata, Y., Aoyama, M., Hirose, K., Sanada, Y., Torii, T., Tsubono, T., Tsunume, D.,
 996 Yamada, M., 2014. Distribution of radionuclides in surface seawater obtained by
 997 an aerial radiological survey. *J. Nucl. Sci. Technol.* 51, 1059-1063.
 998 <https://doi.org/10.1080/00223131.2014.914451>.
- 999 Inomata, Y., Aoyama, M., Tsubono, T., Tsumune, D., Hirose, K., 2016. Spatial and
 1000 temporal distributions of ¹³⁴Cs and ¹³⁷Cs derived from the TEPCO Fukushima
 1001 Daiichi Nuclear Power Plant accident in the North Pacific Ocean by using optimal
 1002 interpolation analysis. *Environ. Sci. Proc. Imp.* 18, 126-36.
 1003 DOI:[10.1039/C5EM00324E](https://doi.org/10.1039/C5EM00324E).
- 1004 Israelsson, S, Knudsen, E, 1986. Effect of radioactive fallout from a nuclear power plant
 1005 accident on electrical parameters. *J. Geophys. Res.* 91, 11909-11910.
 1006 <https://doi.org/10.1029/JD091iD11p11909>.
- 1007 Israelsson, S., Shutte, T., Pislser, E., Lundquist, S., 1987. Increased occurrence of
 1008 lightning flashes in Sweden during 1986. *J. Geophys. Res.* 92, 10996-10998.
 1009 <https://doi.org/10.1029/JD092iD09p10996>.
- 1010 Jana, S., Maitra, A., 2019, Electric field variation in clear and convective conditions at a
 1011 tropical urban location. *J. Geophys. Res. Atmos.* 124, 2068- 2078.
 1012 <https://doi.org/10.1029/2018JD028310>.
- 1013 Kajino, M., Sekiyama, T.T., Igarashi, Y., Katata, G., Sawada, M., Adachi, K., Zaizen, Y.,
 1014 Tsuruta, H., Nakajima, T., 2019. Deposition and Dispersion of Radio-Cesium
 1015 Released Due to the Fukushima Nuclear Accident: Sensitivity to Meteorological
 1016 Models and Physical Modules. *J. Geophys. Res. Atmos.* 124, 1823-1845.
 1017 <https://doi.org/10.1029/2018JD028998>.
- 1018 Kanai, Y., 2012. Monitoring of aerosols in Tsukuba after Fukushima Nuclear Power
 1019 Plant incident in 2011. *J. Environ. Radioact.* 111, 33-37.
 1020 DOI:[10.1016/j.jenvrad.2011.10.011](https://doi.org/10.1016/j.jenvrad.2011.10.011).

1021 Katata, G., Chino, M., Kobayashi, T., Terada, H., Ota, M., Nagai, H., Kajino, M.,
1022 Draxler, R., Hort, M.C., Malo, A., Torii, T., Sanada, Y., 2015. Detailed source
1023 term estimation of the atmospheric release for the Fukushima Daiichi Nuclear
1024 Power Station accident by coupling simulations of an atmospheric dispersion
1025 model with an improved deposition scheme and oceanic dispersion model. *Atmos.*
1026 *Chem. Phys.* 15, 1029-1070. <https://doi.org/10.5194/acp-15-1029-2015>.

1027 Katata, G., Ota, M., Terada, H., Chino, M., Nagai, H., 2012. Atmospheric discharge and
1028 dispersion of radionuclides during the Fukushima Daiichi Nuclear Power Plant
1029 accident. Part I: source term estimation and local-scale atmospheric dispersion in
1030 early phase of the accident. *J. Environ. Radioact.* 109, 103-113.
1031 <https://doi.org/10.1016/j.jenvrad.2012.02.006>.

1032 Kawano, M., Ikebe, Y., Shimo, M., 1969. Measurements of attachment coefficients of
1033 small ions and radioactive ions to condensation nuclei, in: Coroniti, S., Hughes J.
1034 (Eds.), *Planetary Electrodynamics*, Gordon and Breach, New York, 1, pp. 136-
1035 154.

1036 Kita K, Kasahara R, Tanaka M, Sato K, Demizu H, et al., 2013. Variation of atmospheric
1037 concentrations of I-131, Cs-134 and Cs-137 observed over eastern Japan:
1038 contribution of leakage from Fukushima Dai-ichi Nuclear Power Plant and
1039 secondary emission from soil and vegetation. *Geophys. Res. Abstr.* 15, EGU2013-
1040 5090.

1041 Kondo G., 1959. The Recent Status of Secular Variation of the Atmospheric Electric
1042 Elements and their Relation to the Nuclear Explosions, *Memoirs of Kakioka*
1043 *Magnetic Observatory*, 9, 2-6.

1044 Koo, Y.-H., Yang, Y.-S., Song, K.-W., 2014. Radioactivity release from the Fukushima
1045 accident and its consequences: A review. *Prog. Nucl. Energy* 74, 61-70.
1046 <https://doi.org/10.1016/j.pnucene.2014.02.013>.

1047 Korobova, E., Linnik, V., Chizhikova, N., 2008. The history of the Chernobyl ¹³⁷Cs
1048 contamination of the flood plain soils and its relation to physical and chemical
1049 properties of the soil horizons (a case study). *J. Geochem. Explor.* 96, 236-255.
1050 <https://doi.org/10.1016/j.gexplo.2007.04.014>.

1051 Korsakissok, I., Mathieu, A., Didier, D., 2013. Atmospheric dispersion and ground
1052 deposition induced by the Fukushima Nuclear power plant accident: a local-scale
1053 simulation and sensitivity study. *Atmos. Environ.* 70, 267-279.
1054 <https://doi.org/10.1016/j.atmosenv.2013.01.002>.

1055 Kumamoto, Y., Aoyama, M., Hamajima, Y., Nishino, S., Murata, A., Kikuchi, T., 2016.
1056 Meridional distribution of Fukushima-derived radiocesium in surface seawater
1057 along a trans-Pacific line from the Arctic to Antarctic Oceans in summer 2012. *J.*
1058 *Radioanal. Nucl. Chem.* 307, 1703-1710. <https://doi.org/10.1007/s10967-015-4439-0>.

1060 Kumamoto, Y., Aoyama, M., Hamajima, Y., Nishino, S., Murata, A., Kikuchi, T., 2017.
1061 Radiocesium in the western subarctic area of the north Pacific Ocean, Bering Sea,
1062 and Arctic Ocean in 2013 and 2014. *Appl. Radiat. Isot.* 126, 88-92.
1063 <https://doi.org/10.1016/j.apradiso.2017.02.036>.

1064 Kumamoto Y, Aoyama M, Hamajima Y, Nishino S, Murata A, Kikuchi T., 2019.
1065 Radiocesium in the western subarctic area of the North Pacific ocean, Bering Sea,

1066 and Arctic Ocean in 2015 and 2017. *Polar Science* 2019; 21: 228-232.
1067 <https://doi.org/10.1016/j.polar.2018.08.007>.

1068 Leelőssy, Á., Lagzi, I., Kovács, A., Mészáros, R., 2018. A review of numerical models to
1069 predict the atmospheric dispersion of radionuclides. *J. Environ. Radioact.* 182, 20-
1070 33. <https://doi.org/10.1016/j.jenvrad.2017.11.009>.

1071 Lelieveld J., Kunkel D, Lawrence M.G., 2012. Global risk of radioactive fallout after
1072 major nuclear accidents. *Atmos. Chem. Phys.* 12, 4245-4258. DOI:[10.5194/acp-
1073 12-4245-2012](https://doi.org/10.5194/acp-12-4245-2012).

1074 Lin, W., Chen, L., Yu, W., Ma, H., Zeng, Z., Lin, J., Zeng, S., 2015. Radioactivity
1075 impacts of the Fukushima Nuclear Accident on the atmosphere. *Atmos. Environ.*
1076 102, 311-322. <https://doi.org/10.1016/j.atmosenv.2014.11.047>.

1077 Marzo, G.A., 2014. Atmospheric transport and deposition of radionuclides released after
1078 the Fukushima Dai-ichi accident and resulting effective dose. *Atmos. Environ.* 94,
1079 709-722. <https://doi.org/10.1016/j.atmosenv.2014.06.009>.

1080 Mathieu, A., Kajino, M., Korsakissok, I., Périllat, R., Quélo, D., Quérel, A., Saunier, O.,
1081 Sekiyama, T.T., Igarashi, Y., Didier, D., 2018. Fukushima Daiichi-derived
1082 radionuclides in the atmosphere, transport and deposition in Japan: A review.
1083 *Appl. Geochem.* 91, 122-139. <https://doi.org/10.1016/j.apgeochem.2018.01.002>.

1084 MEXT (Ministry of Education, Culture, Sports, Science and Technology), 2007. SPEEDI
1085 System for Prediction of Environmental Emergency Dose Information, Pamphlet
1086 of SPEEDI.

1087 MEXT (Ministry of Education, Culture, Sports, Science and Technology), 2011. Results
1088 of Airborne Monitoring by the Ministry of Education, Culture, Sports, Science
1089 and Technology and the U.S. Department of Energy, Annex 1,
1090 https://radioactivity.nsr.go.jp/en/contents/4000/3180/24/1304797_0506.pdf.

1091 Morino, Y., Ohara, T., Nishizawa, M., 2011. Atmospheric behavior, deposition, and
1092 budget of radioactive materials from the Fukushima Daiichi nuclear power plant
1093 in March 2011. *Geophys. Res. Lett.* 38. <https://doi.org/10.1029/2011GL048689>.

1094 Morino, Y., Ohara, T., Watanabe, M., Hayashi, S., Nishizawa, M., 2013. Episode
1095 analysis of deposition of radiocesium from the Fukushima Daiichi nuclear power
1096 plant accident. *Environ. Sci. Technol.* 47, 2314-2022.
1097 <https://doi.org/10.1021/es304620x>.

1098 Nagai, H., Chino, M., Yamazawa, H., 1999. Development of scheme for predicting
1099 atmospheric dispersion of radionuclides during nuclear emergency by using
1100 atmospheric dynamic model. *J. At. Energy. Soc. Jpn.* 41, 777-785, [in Japanese
1101 with English abstract].

1102 Nakajima, T., Misawa, S., Morino, Y., Tsuruta, H., Goto, D., Uchida, J., Takemura, T.,
1103 Ohara, T., Oura, Y., Ebihara M., Satoh, M., 2017. Model depiction of the
1104 atmospheric flows of radioactive cesium emitted from the Fukushima Daiichi
1105 Nuclear Power Station accident. *Prog. Earth Planet. Sci.* 4, 2.
1106 <https://doi.org/10.1186/s40645-017-0117-x>.

1107 Nicoll, K.A., 2012. Measurements of atmospheric electricity aloft. *Surv. Geophys.* 33,
1108 991-1057. <https://doi.org/10.1007/s10712-012-9188-9>.

1109 Nicoll, K. A., R. G. Harrison, V. Barta, J. Bor, R. Brugge, A. Chillingarian, J. Chum, A.
1110 K. Georgoulas, A. Guha, K. Kourtidis, M. Kubicki, E. Mareev, J. Matthews, H.
1111 Mkrtychyan, A. Odzimek, J.-P. Raulin, D. Robert, H. G. Silva, J. Tacza, Y. Yair,

1112 and R.Yaniv, 2019, A global atmospheric electricity monitoring network for
1113 climate and geophysical research, *J. Atmos. Sol-Terr. Phys.*, 184, 18-29.
1114 <https://doi.org/10.1016/j.jastp.2019.01.003>.

1115 Ogawa T., Morita M., 1977. Electric conductivity, ozon concentration, water vapour
1116 pressure, and electric current at 20-350mb altitude (in Japanese), *Taiki-Denki*
1117 *Kenkyu* (Proceedings of the Society of Atmospheric Electricity of Japan), 18,
1118 107-114, doi:10.11501/9666801, ISSN 0286-3405.
1119 <http://dl.ndl.go.jp/info:ndljp/pid/9666801>.

1120 Panitz, H.-J., Matzerath, C., Päsler-Sauer J., 1989. UFOMOD Atmospheric Dispersion
1121 and Deposition, Institut für Neutronenphysik und Reaktortechnik Projekt LWR-
1122 Sicherheit, KfK 4332, ISSN 0303-4003.

1123 Pierce, E.T., 1959. Some calculations on radioactive fallout with especial reference to the
1124 secular variations in potential gradient at Eskdalemuir, Scotland. *Pure Appl.*
1125 *Geophys.* 42, 145-151. <https://doi.org/10.1007/BF02113404>.

1126 Pierce, E.T., 1972. Radioactive fallout and secular effects in atmospheric electricity. *J.*
1127 *Geophys. Res.* 77, 482-487. <https://doi.org/10.1029/JC077i003p00482>.

1128 Piper, I. M., Bennett. A. J., 2012. Observations of the atmospheric electric field during
1129 two case studies of boundary layer processes. *Environ. Res. Lett.* 7, 014017.
1130 <https://doi.org/10.1088/1748-9326/7/1/014017>.

1131 Povinec, P.P., Gera, M., Holy, K., Hirose, K., Lujaniene, G., Nakano, M., Plastino, W.,
1132 Sykora, I., Bartok, J., Gazak, M., 2013. Dispersion of Fukushima radionuclides in
1133 the global atmosphere and the ocean. *Appl. Radiat. Isot.* 81, 383–392.
1134 <https://doi.org/10.1016/j.apradiso.2013.03.058>.

1135 Pudykiewicz, J., 1988. Numerical simulation of the transport of radioactive cloud from
1136 the Chernobyl nuclear accident. *Tellus B Chem. Phys. Meteorol.* 40, 241-259.
1137 <https://doi.org/10.1111/j.1600-0889.1988.tb00295.x>.

1138 Pullen, J., Chang, J., Hanna, S., 2013. Air-sea transport, dispersion, and fate modeling in
1139 the vicinity of the Fukushima nuclear power plant e a special conference session
1140 summary. *Bull. Am. Meteorol. Soc.* 94, 25-30. <https://doi.org/10.1175/BAMS-D-11-00158.1>.

1142 Quérel, A, Roustan, Y., Quélo, D., Benoit, J.-P., 2015. Hints to discriminate the choice of
1143 wet deposition models applied to an accidental radioactive release. *Int. J. Environ.*
1144 *Pollut.* 56, 268-279. <https://doi.org/10.1504/IJEP.2015.077457>.

1145 Retalis, D., Pitta, A., 1989. Effects on electrical parameters at Athens, Greece by
1146 radioactive fallout from a nuclear power plant accident. *J. Geophys. Res.* 94,
1147 13093–13097, <https://doi.org/10.1029/JD094iD11p13093>.

1148 Retalis, D.A., 1987. Chernobyl Reactor Accident Consequences On Small Atmospheric
1149 Ions Concentration Above Athens. *Pure Appl. Geophys.* 125, 669-678.
1150 <https://doi.org/10.1007/BF00879578>.

1151 Rycroft, M.J., Harrison, R.G., Nicoll, K.A., Mareev, E.A., 2008. An Overview of Earth's
1152 Global Electric Circuit and Atmospheric Conductivity. *Space Sci. Rev.* 137, 83.
1153 <https://doi.org/10.1007/s11214-008-9368-6>.

1154 Rycroft, M.J., Nicoll, K.A., Aplin, K., Harrison, R.G., 2012. Recent advances in global
1155 electric circuit coupling between the space environment and the troposphere. *J.*
1156 *Atmos. Sol-Terr. Phys.* 90-91, 198-211.
1157 <https://doi.org/10.1016/j.jastp.2012.03.015>.

1158 Sanada, Y., Matsunaga, T., Yanase, N., Nagao, S., Amano, H., Takada, H., Tkachenko,
1159 Y., 2002. Accumulation and potential dissolution of Chernobyl-derived
1160 radionuclides in river bottom sediment. *Appl. Radiat. Isot.* 56, 751-760.
1161 [https://doi.org/10.1016/S0969-8043\(01\)00274-3](https://doi.org/10.1016/S0969-8043(01)00274-3).

1162 Saito, K., Shimbori, T., Draxler, R., Hara, T., Toyoda, E., Honda, Y., Nagata, K., Fujita,
1163 T., Sakamoto, M., Kato, T., Kajino, M., Sekiyama, T.T., Tanaka, T.Y., Maki, T.,
1164 Terada, H., Chino, M., Iwasaki, T., Hort M.C., Leadbetter, S.J., Wotawa, G.,
1165 Arnold, D., Maurer, C., Malo, A., Servranckx, R., Chen, P., 2015. Contribution of
1166 JMA to the WMO Technical Task Team on Meteorological Analyses for
1167 Fukushima Daiichi Nuclear Power Plant Accident and Relevant Atmospheric
1168 Transport Modeling at MRI. Technical Reports of the Meteorological Research
1169 Institute, No. 76, Japan Meteorological Agency, DOI : 10.11483/mritechrepo.76.

1170 Sato, Y., Takigawa, M., Sekiyama, T.T., Kajino, M., Terada, H., et al. 2018. Model
1171 Intercomparison of Atmospheric ¹³⁷Cs From the Fukushima Daiichi Nuclear
1172 Power Plant Accident: Simulations Based on Identical Input Data. *J. Geophys.*
1173 *Res. Atmos.* 123, 11, 748-11,765. <https://doi.org/10.1029/2018JD029144>.

1174 SCJ (Science Council of Japan), 2014. A review of the model comparison of
1175 transportation and deposition of radioactive materials released to the environment
1176 as a result of the Tokyo Electric Power Company's Fukushima Daiichi Nuclear
1177 Power Plant accident, Committee on Comprehensive Synthetic Engineering,
1178 pp.103.

1179 Shigeno, N., Takizawa, T., Itoh, N., Yokoyama, M., Owada, T., 2001. Preliminary test
1180 for atmospheric electricity measured using an electrostatic sensor, *Gijutsu*
1181 *Hookoku*, 112, 8-13.

1182 Smith JN, Rossi V, Buessler KO, Cullen JT, Cornett J, Nelson R, Macdonald, A.
1183 M., Robert, M., Kellogg, J., 2017. Recent Transport History of Fukushima
1184 Radioactivity in the Northeast Pacific Ocean. *Environmental Science &*
1185 *Technology*; 51: 10494-10502. DOI: 10.1021/acs.est.7b02712

1186 Stoehlker, U., Nikkinen, M., Gheddou, A., 2011. Detection of radionuclides emitted
1187 during the Fukushima nuclear accident with the CTBT radionuclide network, in:
1188 Proceedings of the 2011 Monitoring Research Review: Ground-Based Nuclear
1189 Explosion Monitoring, 715-724. [https://www.osti.gov/biblio/1027453-](https://www.osti.gov/biblio/1027453-proceedings-monitoring-research-review-ground-based-nuclear-explosion-monitoring-technologies)
1190 [proceedings-monitoring-research-review-ground-based-nuclear-explosion-](https://www.osti.gov/biblio/1027453-proceedings-monitoring-research-review-ground-based-nuclear-explosion-monitoring-technologies)
1191 [monitoring-technologies](https://www.osti.gov/biblio/1027453-proceedings-monitoring-research-review-ground-based-nuclear-explosion-monitoring-technologies).

1192 Stohl, A., Seibert, P., Wotawa, G., Arnold, D., Burkhart, J.F., Eckhardt, S., Tapia, C.,
1193 Vargas, A., Yasunari, T.J., 2012. Xenon-133 and caesium-137 releases into the
1194 atmosphere from the Fukushima Dai-ichi nuclear power plant: determination of
1195 the source term, atmospheric dispersion, and deposition. *Atmos. Chem. Phys.* 12,
1196 2313-2343. <https://doi.org/10.5194/acp-12-2313-2012>.

1197 Takeda, M., Yamauchi, M., Makino, M., Owada, T., 2011. Initial effect of the Fukushima
1198 accident on atmospheric electricity. *Geophys. Res. Lett.* 38, L15811,
1199 <https://doi.org/10.1029/2011GL048511>.

1200 Talerko, N., 2005a. Mesoscale modelling of radioactive contamination formation in
1201 Ukraine caused by the Chernobyl accident. *J. Environ. Radioact.* 78, 311-
1202 329. <https://doi.org/10.1016/j.jenvrad.2004.04.008>.

1203 Talerko, N., 2005b. Reconstruction of ^{131}I radioactive contamination in Ukraine caused
1204 by the Chernobyl accident using atmospheric transport modeling. *J. Environ.*
1205 *Radioact.* 84, 343-362. <https://doi.org/10.1016/j.jenvrad.2005.04.005>.

1206 Terada, H., Chino, M., 2008a. Development of an atmospheric dispersion model for
1207 accidental discharge of radionuclides with the function of simultaneous prediction
1208 for multiple domains and its evaluation by application to the Chernobyl nuclear
1209 accident. *J. Nucl. Sci. Technol.* 45, 920–931. [10.1080/18811248.2008.9711493](https://doi.org/10.1080/18811248.2008.9711493).

1210 Terada, H., Katata, G., Chino, M., Nagai, H., 2012. Atmospheric discharge and
1211 dispersion of radionuclides during the Fukushima Dai-ichi Nuclear Power Plant
1212 accident. Part II: verification of the source term and analysis of regional-scale
1213 atmospheric dispersion. *J. Environ. Radioact.* 112, 141–154.
1214 <https://doi.org/10.1016/j.jenvrad.2012.05.023>.

1215 Terada, H., Nagai, H., Furuno, A., Kakefuda, T., Harayama, T., Chino, M., 2008b.
1216 Development of worldwide version of system for prediction of environmental
1217 emergency dose information: WSPEEDI 2nd version. *Trans. At. Energy Soc. Jpn.*
1218 7, 257–267, [in Japanese with English abstract].
1219 <https://doi.org/10.3327/taesj.J07.045>.

1220 Thakur, P., Ballard, S., Nelson, R., 2013. An overview of Fukushima radionuclides
1221 measured in the northern hemisphere. *Sci. Total Environ.* 458-460, 577-613.
1222 DOI:[10.1016/j.scitotenv.2013.03.105](https://doi.org/10.1016/j.scitotenv.2013.03.105).

1223 Tsubono, T., Misumi, K., Tsumune, D., Bryan, F.O., Hirose, K., Aoyama, M., 2016.
1224 Evaluation of radioactive cesium impact from atmospheric deposition and direct
1225 release fluxes into the North Pacific from the Fukushima Daiichi nuclear power
1226 plant. *Deep Sea Res. Part 1 Oceanogr. Res. Pap.* 115, 10-21.
1227 <https://doi.org/10.1016/j.dsr.2016.02.019>.

1228 Tsumune, D., Tsubono, T., Aoyama, M., Hirose, K., 2012. Distribution of oceanic ^{137}Cs
1229 from the Fukushima Dai-ichi Nuclear Power Plant simulated numerically by a
1230 regional ocean model. *J. Environ. Radioact.* 111, 100-08.
1231 <https://doi.org/10.1016/j.jenvrad.2011.10.007>.

1232 Tuomi T.J., 1982. The atmospheric electrode effect over snow. *J. Atmos. Sol. Terr. Phys.*
1233 44, 737-745. [https://doi.org/10.1016/0021-9169\(82\)90002-2](https://doi.org/10.1016/0021-9169(82)90002-2).

1234 Tuomi T.J., 1989. Ten Year Summary 1977-1986 of Atmospheric Electricity Measured at
1235 Helsinki-Vantaa Airport, Finland. *Geophysica* 25, 1-20.

1236 Tuomi, T.J., 1988. Observations of atmospheric electricity 1986. *Geophys. Publ.* 7,
1237 551.506.1, Finnish Meteorological Institute, Helsinki, 61 pp.

1238 Uman, M.A., 1986. All about lightning, 173 pp, ISBN 0-486-25237-X, Dover Pub. Inc.,
1239 Mineola, USA.

1240 UNSCEAR (United Nations Scientific Committee on the Effect of Atomic Radiation),
1241 2000. Report to the General Assembly with Scientific Annexes, Volume II, Annex
1242 J - Exposures and effects of the Chernobyl accident. UNSCEAR, United Nations,
1243 New York.

1244 UNSCEAR (United Nations Scientific Committee on the Effect of Atomic Radiation),
1245 2013. Report to the General Assembly with Scientific Annexes, Volume I, Annex
1246 A - Levels and effects of Radiation Exposure due to the nuclear accident after the
1247 2011 great east - Japan earthquake and tsunami, UNSCEAR, United Nations, New
1248 York.

1249 Warzecha, S., 1987. Results of atmospheric electricity measurements at Swider after the
1250 Chernobyl nuclear power plant accident. *Publs. Inst. Geophys. Pol. Acad. Sc.*, D-
1251 26 (198).

1252 Warzecha, S., 1991. Long-lasting radioactive contamination at Swider and its effect on
1253 the atmospheric electricity parameters after the Chernobyl nuclear power plant
1254 accident. *Publs. Inst. Geophys. Pol. Acad. Sc.*, D-35 (238).

1255 Williams, E., Mareev, E., 2014. Recent progress on the global electrical circuit. *Atmos.*
1256 *Res.* 135-136, 208-227. <http://doi.org/10.1016/j.atmosres.2013.05.015>.

1257 Williams, E.R., 2009. The global electrical circuit: A review. *Atmos. Res.* 91, 140-152.
1258 <http://doi.org/10.1016/j.atmosres.2008.05.018>.

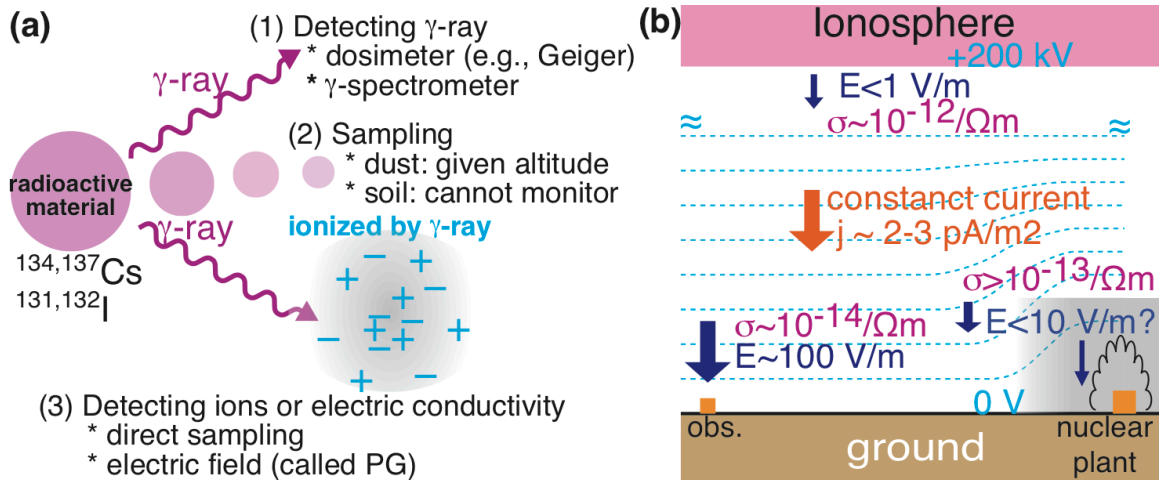
1259 Winiarek, V., Bocquet, M., Saunier, O., Mathieu, A., 2012. Estimation of errors in the
1260 inverse modeling of accidental release of atmospheric pollutant: application to the
1261 reconstruction of the cesium-137 and iodine-131 source terms from the
1262 Fukushima Daiichi power plant. *J. Geophys. Res.* 117, D05122.
1263 <https://doi.org/10.1029/2011JD016932>.

1264 WMO (World Meteorological Organization), 2013. The World Meteorological
1265 Organization's Evaluation of Meteorological Analyses for the Radionuclide
1266 Dispersion and Deposition from the Fukushima Daiichi Nuclear Power Plant
1267 Accident. Report on the third meeting of the WMO Task Team on meteorological
1268 analyses for the Fukushima-Daiichi nuclear power plant accident, annex III,
1269 [http://www.wmo.int/pages/prog/www/CBS-
1270 Reports/documents/WMO_fnpp_final_AnnexIII_4Feb2013_REVISIED_17June20
1271 13.pdf](http://www.wmo.int/pages/prog/www/CBS-Reports/documents/WMO_fnpp_final_AnnexIII_4Feb2013_REVISIED_17June2013.pdf).

1272 Yamauchi, M., 2012. Secondary wind transport of radioactive materials after the
1273 Fukushima accident. *Earth Planets Space* 64, e1-e4.
1274 <https://doi.org/10.5047/eps.2012.01.002>.

1275 Yamauchi, M., Takeda, M., Makino, M., Owada, T., Miyagi, I., 2012. Settlement process
1276 of radioactive dust to the ground inferred by the atmospheric electric field
1277 measurement. *Ann. Geophys.* 30, 49-56. [https://doi.org/10.5194/angeo-30-49-
1278 2012](https://doi.org/10.5194/angeo-30-49-2012).

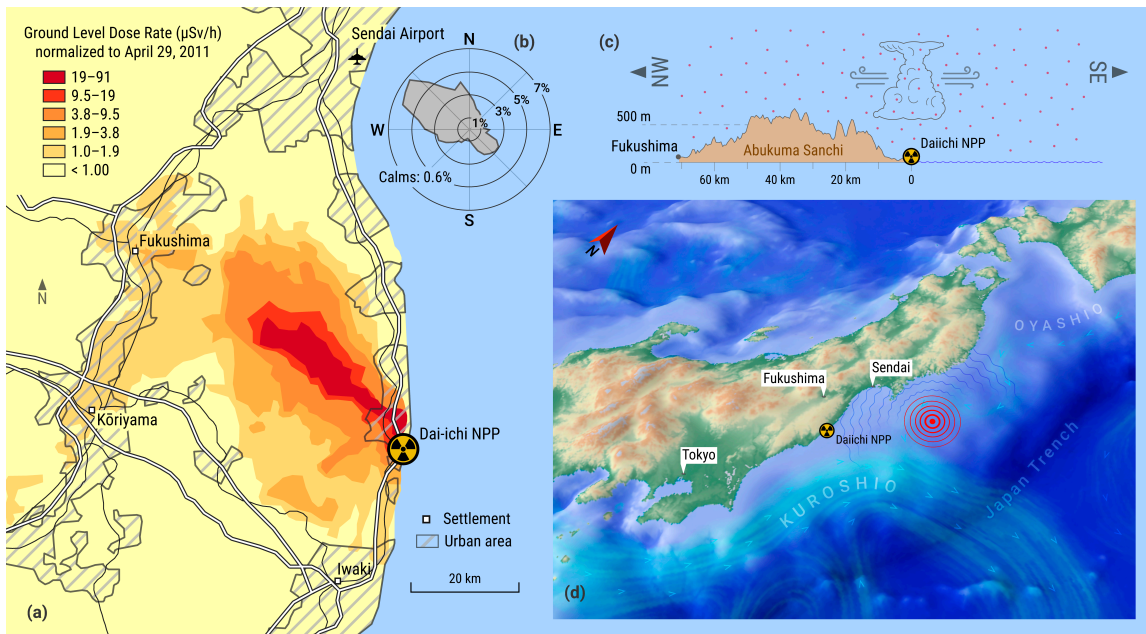
1279 Yamauchi, M., Takeda, M., Nagamachi, S., 2018. Effect of enhanced ionizing radiation
1280 on the cloud electricity after the Fukushima nuclear accident. *Earth Planets Space*
1281 70, 48. <https://doi.org/10.1186/s40623-018-0780-1>.
1282
1283



1284

1285 Figure 1. Illustrations of the monitoring principle of enhanced ionized radiation in the
 1286 atmosphere: (a) Instead of direct detection of floating radioactive materials or ionizing
 1287 radiation, one can also monitor the resultant ion density or even the vertical electric field
 1288 (Potential Gradient: PG), (b) Local PG decreases (distance between equipotential surface
 1289 increases) if the local ionospheric conductivity increases because the background electric
 1290 field that is maintained on a global scale between the ionosphere and the ground. The PG
 1291 works under fair weather when electric field due to free electric charges (such as
 1292 electrified cloud) is not significant compared to background electric field.

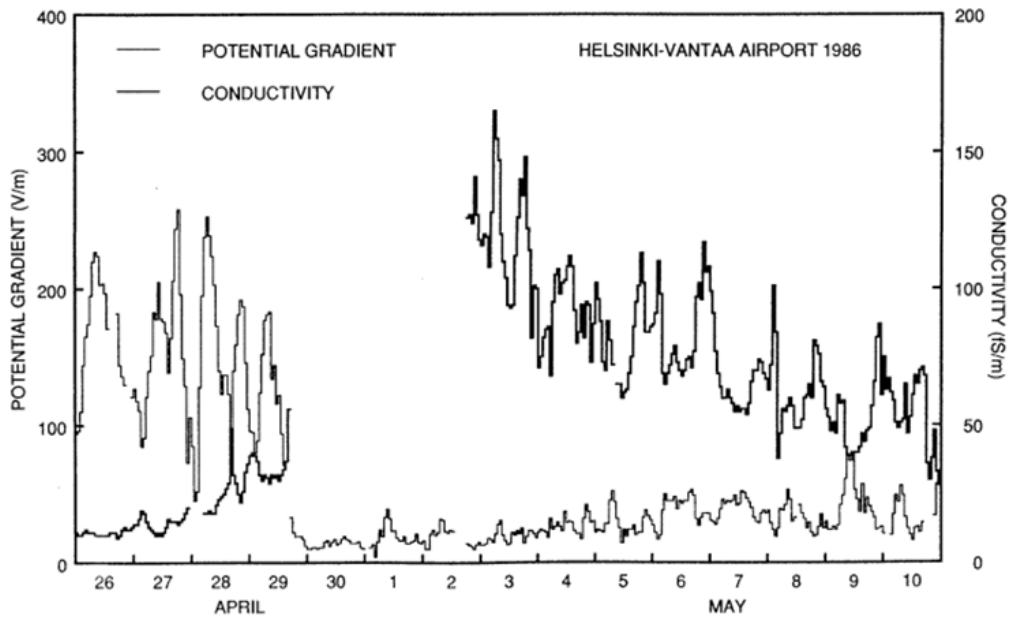
1293



1294

1295 Figure 2. (a) Dose measurement map based on airborne monitoring map from April 6 to
 1296 29 by MEXT and DOE (Readings of air dose rate 1 m from the ground monitored inside
 1297 80 km zone of FNPP) (MEXT, 2011), (b) A wind rose of Sendai Airport (situated about
 1298 90 km from FNPP), the seven-year March summary (March 1, 2012 – March 31, 2018)
 1299 (IEM), (c) The terrain profile between FNPP and the city of Fukushima with an Abukuma
 1300 Sanchi Mountain situated between them and a schematic diagram of a dominant
 1301 radionuclides release to the NW direction (to inland areas, toward the city of Fukushima)
 1302 and the SE direction (to the ocean), (d) The 3D map showing topography and bathymetry
 1303 near FNPP and epicenter of the earthquake. The center of concentric circles represents
 1304 the epicenter of the earthquake which occurred near the Japan Trench, a line where the
 1305 Pacific plate is subducting under the North American plate. Turquoise flow lines and
 1306 arrows represent ocean currents.

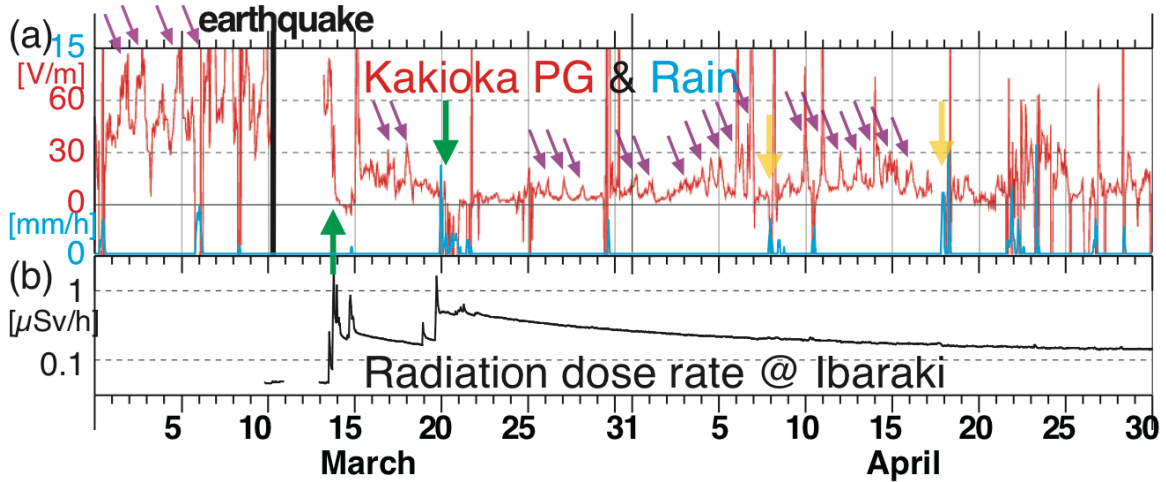
1307



1308

1309 Figure 3. One-hour resolution PG (solid line) and atmospheric conductivity (bold line) at
 1310 Helsinki for one month in April-May 1986, when the Chernobyl nuclear accident took
 1311 place (after Tuomi, 1988). Rain-induced fallout took place on April 29, 1986.

1312

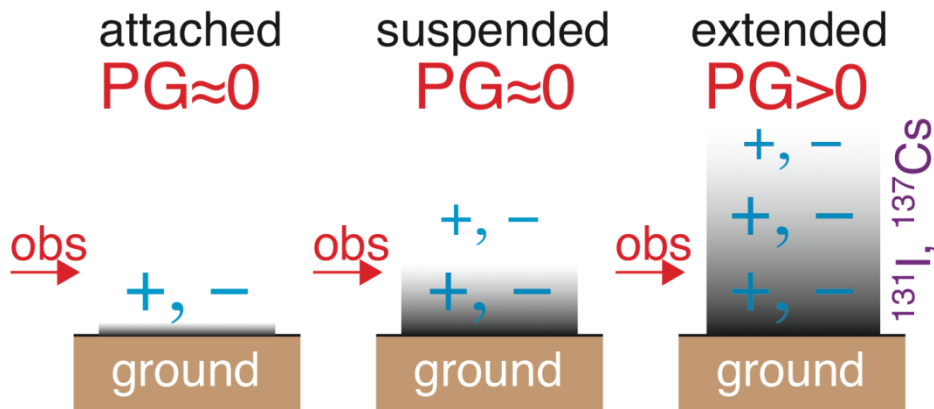


1313

1314 Figure 4. (a) Hourly averaged PG (red line) and one-hour accumulated rainfall (blue line)
 1315 at Kakioka. (b) Hourly averaged radiation dose rate at Ibaraki-cho (nearest station of
 1316 radiation observation to Kakioka). The time resolutions of raw data are 1 second for PG
 1317 data and 10 minutes for the rainfall and radiation dose rate. The data gap from March 11
 1318 to ~March 14 is due to the power failure caused by the earthquake. The purple arrows in
 1319 the upper panel indicate variation due to diurnal wind under fair weather, with peak local
 1320 time (LT) near noon after the accident whereas it is 3-4 LT before the accident
 1321 (Yamauchi et al., 2012).

1322

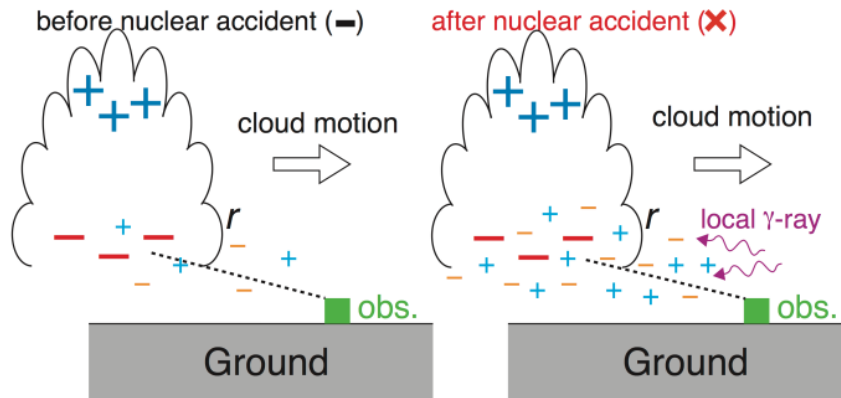
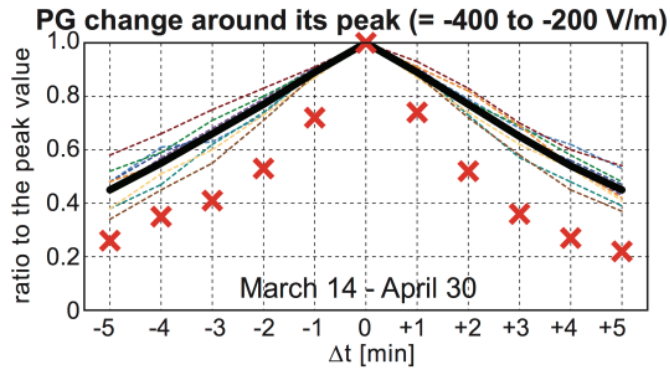
1323



1324

1325 Figure 5. Expected PG for different degrees of contamination: firmly-attached (or
 1326 migrated), loosely floating (or suspended near the ground), and re-suspended to extended
 1327 altitude (Yamauchi et al., 2012).

1328



1329

1330 Figure 6. Relative PG values compared to peak values during 10 min centered at the peak
 1331 time (top) and illustration of electrostatic shielding of cloud charge by surrounding ions
 1332 (bottom) (Yamauchi et al., 2018). Red X marks are for 2011, dashed lines are for the
 1333 other years, and the thick line is for the average. The shielding is larger for longer
 1334 distances (with more shielding charges).